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Reverse tracing of short-term earthquake precursors

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12 Abstract

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We introduce a new approach to short-term earthquake prediction named "*Reverse Tracing of Precursors*" (RTP), since it considers precursors in reverse order of their appearance. First, we detect the "candidates" for the short-term precursors; in our case, these are newly introduced chains of earthquakes reflecting the rise of an earthquake correlation range. Then we consider each chain, one by one, checking whether it was preceded by an intermediate-term precursor in its vicinity. If *yes*, we regard this chain as a precursor; in prediction it would start a short-term alarm. The chain indicates the narrow area of possibly complex shape, where an intermediate-term precursor should be looked for. This makes possible to detect precursors undetectable by the direct analysis.

RTP can best be described on an example of its application; we describe retrospective prediction of two prominent Californian earthquakes—Landers (1992), M = 7.6, and Hector Mine (1999), M = 7.3, and suggest a hypothetical prediction algorithm. Its validation is considered in subsequent studies, starting from [Shebalin et al., Phys. Earth Planet. Int., in press]. The goal of this paper is to describe RTP methodology, since it has potentially important applications to many other data and to prediction of other critical phenomena besides earthquakes. In particular, it might vindicate some short-term precursors, previously

25 rejected as giving many false alarms.

Validation of the algorithm per se requires its application in different regions with a substantial number of strong earthquakes. First (and positive) results are obtained for 21 more strong earthquakes in California ($M \ge 6.4$), Japan ($M \ge 7.0$) and the Eastern Mediterranean ($M \ge 6.5$); these results are described elsewhere. The final validation requires, as always, prediction

in advance for which this study sets up a base. We have the first case of a precursory chain reported in advance of a subsequent

- 30 strong earthquake (Tokachi-oki, near Hokkaido island, Japan), 25 September 2003, M = 8.1.
- 31 Possible mechanisms underlying RTP are outlined.

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33 Keywords: Short-term earthquake prediction; Long-range correlation; Earthquake chains; Precursors; Instability

34

 1. Introduction
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 1.1. Generation of strong earthquakes—a
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 non-localized process
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 Seismicity is commonly recognized as a part of the
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geodynamics (Aki, 2003; Bird, 1998; Keilis-Borok,

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1990; King et al., 2002; Press, 1965; Rundquist and 40 Soloviev, 1999; Scholz, 1990); in seismically active ar-41 eas the earthquakes accommodate a considerable frac-42 tion of tectonic development of the lithosphere. That 43 development goes on in multiple time-, space-, and 44 energy-scales and preparation of strong earthquakes 45 46 is not an exception. Accordingly, while the target of earthquake prediction-a strong earthquake-is a lo-47 calized event, the process of its generation is not local-48 ized. Strictly speaking, its time scales range from ge-49 ological to seconds in time, and spatial scales-from 50 global to microscopic (Turcotte, 1997; Keilis-Borok, 51 1990); however, in prediction research a truncated 52 scaling is usually considered: from tens of years to 53 days, and from hundreds of kilometers to kilometer. 54

This multiplicity of scales is reflected in the gen-55 eral concept of the seismically active lithosphere as 56 a hierarchical dissipative non-linear system, persis-57 tently self-organizing from time to time into the criti-58 cal phenomena-the strong earthquakes (Blanter and 59 Shnirman, 1997; Bowman et al., 1998; Gabrielov et al., 60 61 1994, 2000; Jaume and Sykes, 1999; Keilis-Borok, 1990; Rundle et al., 2000; Sornette, 2000; Turcotte, 62 1997; Zaliapin et al., 2002a). Among manifestations 63 of that selforganization are premonitory seismicity 64 patterns-the spatio-temporal patterns of seismic-65 ity emerging as a strong earthquake approaches 66 (Aki, 2003; Buffe and Varnes, 1993; Caputo et al., 67 1983; Gabrielov and Newman, 1994; Jin et al., 68 2003;Keilis-Borok, 1990, 1996, 2000; Keilis-Borok 69 et al., 1990a,b, 1964, 1999, 2002; Knopoff et al., 70 1996; Kossobokov et al., 1995, 2003; Ma et al., 1990; 71 72 Mogi, 1985; Newman et al., 1995; Novikova et al., 2002; Press, 1965; Press and Allen, 1995; Romanow-73 icz, 1993; Rotwain and Novikova, 1999; Shebalin 74 et al., 2000; Turcotte, 1997; Zaliapin et al., 2002a,b, 75 2003; Zöller et al., 2001). A multitude of such pat-76 terns have been reported in rather different scales. 77 Systematically tested are the intermediate-term pat-78 terns (with characteristic lead time of years). Here, 79 we suggest a method to detect the short-term patterns, 80 which have the lead time of months. 81

82 1.2. Reverse Tracing of Precursors (RTP)

We consider the short-term patterns in conjunction
with intermediate-term ones. This is done by RTP
analysis, in which these patterns are detected in the re-

verse order of their appearance: short-term patterns are analyzed first, although they emerge later. Our findings can best be described on a specific example of data analysis.

We describe detection of short-term patterns be-91 fore two prominent Californian earthquakes-Landers 92 (1992), M = 7.6, and Hector Mine (1999), M = 7.3. 93 These are the largest Californian earthquakes since 94 1965-the period, when the earthquake catalog is suf-95 ficiently complete for our analysis. Territory consid-96 ered is shown in Fig. 1. The earthquake catalog is 97 taken from (ANSS/CNSS and NEIC). 98

2. Chains

Our point of departure is provided by the short-term 100 patterns Roc and Accord capturing a premonitory 101 increase in earthquake correlation range. They were 102 found first in models (Gabrielov et al., 2000) and then 103 in observations (Keilis-Borok et al., 2002; Shebalin 104 et al., 2000; Novikova et al., 2002). Other patterns 105 capturing that phenomenon are suggested in Zöller 106 et al. (2001) and Zaliapin et al. (2002b). Here, we 107 introduce the pattern *chain* which is a generalization 108 of Roc and Accord. Qualitatively, a chain is a rapidly 109 extended sequence of small earthquakes that follow 110 each other closely in time and space. 111

- 2.1. Definitions 112
- 2.1.1. Earthquake catalog 113 As in most premonitory patterns of that family 114 (Keilis-Borok, 1996; Kossobokov and Shebalin, 2003) 115 aftershocks are eliminated from the catalog; however, 116 an integral measure of aftershocks sequence b is re-117 tained for each remaining earthquake (main shocks 118 and foreshocks). We use a common representation of 119 the earthquake catalog $\{t_i, \varphi_i, \lambda_i, h_i, m_i, b_i\}, j =$ 120 1, 2, Here, t_j is the time of an earthquake, $t_j \ge$ 121 t_{j-1} ; φ_j and λ_j , latitude and longitude of its epicenter; 122 h_i , focal depth; and m_i , magnitude. We consider the 123 earthquakes with magnitude $m = m_{\min}$. Focal depth 124 is not used in this study. 125

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Fig. 1. Territory considered. Stars mark large earthquakes, targeted for prediction. Dots show background seismicity for the time considered (1965–2003): epicenters of earthquakes with magnitude $m \ge 3$ with aftershocks eliminated. Dashed line is used for time–distance projection of epicenters (Fig. 3 below).

126 2.1.2. Chain

Let us call two earthquakes "neighbors" if: (i) their 127 epicenters lie within a distance r; and (ii) the time in-128 terval between them does not exceed a threshold τ_0 . 129 A chain is the sequence of earthquakes connected by 130 the following rule: each earthquake has at least one 131 neighbor in that sequence; and does not have neigh-132 bors outside the sequence. The threshold r is normal-133 ized by the average distance between the earthquakes 134 with lowest magnitude m in a pair considered. We use 135 a coarse normalization $r = r_0 10^{c\underline{m}}$, c being a numer-136 ical parameter. 137

- 138 Let k be the number of earthquakes thus connected 139 and l—the maximal distance between their epicenters.
- 140 We look for precursors only among the chains with

 $k \ge k_0$ and $l \ge l_0$. These thresholds ensure that our 141 chains are exceptional phenomena. 142

2.1.3. Chain's vicinity 143

To compare location of a chain with locations of 144 strong earthquakes we consider its R-vicinity for-145 mally defined as the union of circles of the radius R146 centered at the epicenters of the chains forming the 147 chain. To smooth the borders of that area we add the 148 dense sequence of circles along the lines connect-149 ing each epicenter in the chain with the two closest 150 ones. The envelope of all the circles is the border of 151 *R*-vicinity of the chain; it is similar to the "Wiener 152 sausage", introduced by N. Wiener in the theory of 153 probability. 154

Table 1

		U				
m _{min}	<i>r</i> ⁰ (km)	с	τ_0 (days)	k_0	<i>l</i> ₀ (km)	<i>R</i> (km)
3.3	50	0.35	20	8	350	75
NT / /·				0.1		

Notations are given in the text, Section 2.1.

155 2.2. Data analysis

We detected the chains defined above using numer-156 ical parameters listed in Table 1. Aftershocks have 157 been identified by a coarse windowing, as described 158 in (Keilis-Borok et al., 2002). The remaining cata-159 log contains 3940 earthquakes. We have found among 160 them nine chains, altogether containing 116 earth-161 quakes: this shows that our chains are indeed excep-162 tional phenomena. Maps of the chains are shown in 163 Fig. 2; shaded areas are their vicinities, defined above. 164 Vital characteristics of each chain are given in Table 2. 165 166 Fig. 3 juxtaposes the chains and strong earthquakes 167 on the time-distance plane; distance is counted along the dashed line shown in Figs. 1 and 2. 168

As we see in Fig. 2 (two panels in the bottom 169 row) and Fig. 3, only the two last chains (#8 and 170 #9) might be regarded as the local short-time pre-171 cursors to the Landers and Hector Mine earthquakes: 172 short-term-because they emerge with the short-term 173 lead times (respectively, 1.7 and 4.6 months); and 174 local-because the target earthquakes occur in their 175 vicinities. However, the other seven chains, if used as 176 precursors, would give false alarms. To reduce their 177 number we introduce the RTP analysis. 178

179 3. Precursory chains

180 3.1. Hypothesis

We hypothesize that a precursory chain (as opposed 181 to a chain giving a false alarm) is preceded by the local 182 intermediate-term precursors formed in the chain's 183 184 *R-vicinity*. This vicinity is not known, until the chain is formed, and its shape might be rather complicated 185 (see Fig. 2). To overcome that impasse we introduce 186 the two-step RTP analysis schematically illustrated in 187 Fig. 4. 188

(i) Search for the chains and determination of their
 R vicinities (Section 2). Each chain is regarded as
 a "candidate" for a short-term precursor.

(ii) Search for the local intermediate-term patterns in 192 the R-vicinities of each chain. They are looked for 193 within T years before the chain; T is an adjustable 194 numerical parameter. If (and only if) such patterns 195 are detected, we regard this chain as a short-term 196 precursor; in prediction it would start a short-term 197 alarm. 198

To complete that description we have to specify 199 intermediate-term patterns used at the second step. 200

We use the *pattern* Σ which reflects premon-202 itory rise of seismic activity. This pattern, intro-203 duced in Keilis-Borok and Malinovskaya (1964), 204 is successfully used in different prediction algo-205 rithms, alone or in combination with other patterns 206 (Keilis-Borok, 1990, 1996, 2000; Keilis-Borok et al., 207 1999, 2002; Kossobokov et al., 1995, 2003; Rotwain 208 and Novikova, 1999). It is defined as a premonitory 209 increase of the total area of the earthquake sources. 210 Emergence of this pattern is captured by the function 211 $\Sigma(t)$ defined in a sliding time-window (Keilis-Borok 212 and Malinovskaya, 1964): 213

$$\sum \left(\frac{T}{s,B}\right) = \sum_{i} 10^{Bm_i}, \ m_i \ge m_{\min}; \ t-s < t_i \le t$$
214

Summation is taken over all main shocks within the 215 time window (t-s, t) in the *R*-vicinity of the chain. We 216 take $B \sim 1$, so that the sum is coarsely proportional 217 to the total area of the fault breaks in the earthquakes' 218 sources (Keilis-Borok, 2002); with B = 0 this sum is 219 the number of earthquakes, with B = 3/2 it is propor-220 tional to their total energy. The emergence of pattern 221 Σ is identified by condition $\Sigma(t) > \Sigma_0$; this thresh-222 old depends on the magnitude of target earthquakes. 223 In previous applications cited above pattern Σ was 224 used as non-local one. We renormalize its numerical 225 parameters to make it local. 226

We detected precursory chains and determined their 228 *R*-vicinities (Section 2). In each vicinity we computed 229 the function $\Sigma(t)$ within time interval T = 5 years and 230 summation interval s = 6 months. We identified as 231 precursory three chains preceded by largest peaks of 232



Fig. 2. Maps of the chains. Detected chains are shown in separate boxes. Circles show epicenters of earthquakes in a chain; their size is proportional to magnitude. The shadowed areas show R-vicinities of the chains. Dates of the beginning and the end of a chain are given at the top of each box. Three chains (1977, 1992, and 1999) shown in bold are identified as precursory ones. The first chain gives a false alarm; two other chains are followed within few months by target earthquakes, Landers and Hector Mine. Other notations are the same as in Fig. 1.

233 $\Sigma(t)$; they can be recognized with the threshold $\Sigma_0 =$ 234 10^{6.7}. Table 2 shows these chains in bold. As we see, 235 identification of the first chain, in 1977, is wrong; in 236 prediction it would give a false alarm. Identification 237 of two other chains, in 1992 and 1999, is correct; each 238 is followed by a target earthquake within few months. 239 The same chains would the selected with the tenfold 230 ameller time interval. T = 6 months. The correct

smaller time interval, T = 6 months. The correspond-

ing threshold is $\Sigma_0 = 10^{5.4}$; it is smaller since smaller 241 number of earthquakes is included in summation. 242

3.4. Hypothetical prediction algorithm

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It remains to define alarms triggered by that precursor. This is a final step in transition from a precursor to algorithmic prediction. We adapt the standard 246

Table 2			
Characteristics	of	the	chain

#	Start	End	Duration (days)	Lead time (months)	Distance from a strong earthquake (km)	Number of earthquakes, k	Maximal distance, <i>l</i> (km)	Largest magnitude	Area of the <i>R</i> -vicinity, $\times 10^3$ (km ²)
28.06	5.1992: Lande	rs earthquake	, M = 7.6						
1	16.07.1969	03.10.1969	80			17	499	5.3	150
2	15.10.1969	19.11.1969	35			12	485	5.6	113
3	26.08.1973	17.10.1973	53			13	381	4.5	150
4	03.06.1977	01.08.1977	60			11	377	4.7	104
5	07.09.1984	26.10.1984	49			9	408	4.6	90
6	08.07.1986	20.07.1986	12			10	543	5.9	122
7	24.12.1989	04.02.1990	41			8	373	5.7	101
8	27.03.1992	08.05.1992	42	1.7	29	17	635	6.1	161
16.10	16.10.1999: Hector Mine earthquake, $M = 7.4$								
9	19.02.1999	01.06.1999	102	4.6	60	11	380	4.9	98

Chains recognized as "precursory" by RTP analysis (Section 3) are shown in bold. Chain #4 would trigger in prediction a false alarm, Chains #8 and #9 would trigger correct alarms.

general scheme of prediction algorithms, widely used
in intermediate-term earthquakes prediction and many
other problems (Keilis-Borok, 2002; Kossobokov and

250 Carlson, 1995, and references therein).

- (i) Prediction is targeted at the main shocks with magnitude M or more; usually the magnitude intervals (M, M + 1) are considered separately.
- (ii) When a precursory chain is detected, a short-term alarm is triggered. It predicts a target earthquake in *R*-vicinity of the chain, within time interval $(t_e, t_e + \tau)$; here t_e is the moment when chain emerged, τ a numerical parameter (duration of alarm). Results of the data analysis suggest to take $\tau = 6$ months.

Possible outcomes of such prediction are illustrated
in Fig. 5. Probabilistic component of prediction is represented by the total time–space covered by alarms
and probabilities of false alarms and failures to predict
(Molchan, 2003).

266 4. Discussion

267 4.1. Summary

This paper introduces RTP analysis in the study of selforganization of seismicity, culminated by a strong earthquake. Precursors with different lead times are considered in reverse order of their appearance. First,

we detect the candidates for short-term precursors: in 272 our case, those are the chains of small earthquakes 273 capturing the rise of earthquake correlation range. A 274 chain determines its narrow vicinity where we look 275 for the local intermediate-term precursor(s), pattern 276 Σ in our case. Its presence in turn indicates the 277 precursory chains. We describe RTP on an exam-278 ple: detecting precursory chains months before two 279 prominent California earthquakes, Landers (1992) 280 and Hector Mine (1999), well isolated in time and 281 space from other comparable earthquakes in that 282 region. 283

4.2. Methodological advantage of RTP

The opposite (direct) analysis would start with trac-285 ing of the intermediate-term patterns hidden in the 286 background seismicity. Almost all of them, known so 287 far, are not local, pattern Σ included. They emerge in 288 the areas whose linear size is up to 10 times larger than 289 the source of the incipient target earthquake (Bowman 290 et al., 1998; Keilis-Borok and Soloviev, 2003); some 291 patterns-even up to 100 times larger (Press and Allen, 292 1995; Romanowicz, 1993). We have found pattern Σ 293 that became local after renormalization: it emerges in 294 the same narrow area (R-vicinity of the chain), where 295 epicenter of a target earthquake lies. As we see in 296 Fig. 2, the shape of that area might be rather com-297 plex, and its size—diverse. To find this area by trying 298 different shapes, sizes, and locations is not realistic. 299

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1965 1970 1975 1980 1985 1990 1995 2000 distance along San-Andreas system, km False alarm Mendocino 0 0 000 8 San-Andreas • 00000 Hector Mine Landers 0 1992 1999 1000 1000 • 500 500 0000 Hector Mine Landers • • \star * . 0 0 1992 1000 500 .ande 0 April May June July

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Fig. 3. Chains and strong earthquakes on the time-distance plain. Distance is counted along the dashed line shown in Fig. 1. Filled and open circles show the chains identified, respectively, as precursory and non-precursory. Other notations are the same as in Fig. 1.

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Fig. 4. Schematic illustration of the *Reverse Tracing of Precursors* (RTP). (Top) Map showing precursory chain and the source of the target earthquake (black). (Bottom) Scheme of analysis in time–space projection. Circles show epicenters forming the chain (dark gray) and preceding it (light gray). The "*R*-vicinity" of the chain is shown in light gray. Star is projection of the epicenter of the target earthquake. The gray rectangle before the chain shows the time–space where rise of activity (pattern Σ) is looked for. White area shows the time–space where this pattern was found; its presence indicates a precursory chain. The chain is detected first, although it emerges after the pattern Σ . Note how a narrow chain determines a much larger time interval where a pattern Σ is looked for. Dark gray area shows the time–space covered by an alarm: within ô months after precursory chain a target earthquake is expected in its *R*-vicinity.

Reverse analysis resolves this impasse, indicating alimited number of chains to consider.

302 4.3. Physical interpretation

RTP seems to be a promising general approach to 303 prediction of critical transitions in non-linear systems: 304 it identifies a rare small-scale phenomenon that car-305 ries a memory of the larger scale history of the sys-306 tem. At the same time, this approach has a natural 307 earth-specific explanation: it follows from the concept 308 that strong earthquake is a result of a lasting large-scale 309 process whose different stages involve different parts 310 of the fault network. Earthquakes in the chain mark 311 the part of the fault network that has started to move in 312 unison months before a target earthquake. Pattern Σ 313 indicates that this synchronization started much ear-314 lier, albeit expressed in a more subtle form. A similar 315 316 step-by-step escalation of instability was observed in

direct analysis: by algorithms M8&MSc (Kossobokov 317 and Shebalin, 2003), and by some other algorithms 318 (Aki, 2003; Shebalin et al., 2000; Keilis-Borok and 319 Soloviev, 2003). 320

Both the chains and the peaks of Σ are sporadic 321 short-lived phenomena not necessarily reflecting the 322 steady trends of seismicity. This is typical for all pre-323 monitory patterns of that family (Keilis-Borok, 2002; 324 Kossobokov and Shebalin, 2003). Probably, both pat-325 terns are the symptoms but not causes of a strong 326 earthquake: they signal its approach but do not trigger 327 it. Similarly sporadic are many observed precursors 328 to other critical phenomena, e.g. economic recessions 329 (Keilis-Borok et al., 2000). 330

4.4. Implications for earthquake prediction 331

- We have applied RTP analysis to target earth-332 quakes of more diverse magnitudes in California 333 and two other regions, Japan and E. Mediter-334 ranean, normalizing the parameters of the algo-335 rithm and considering all known (eight) major 336 types of intermediate-term patterns (Keilis-Borok 337 and Soloviev, 2003). We have first two earthquakes 338 predicted in advance: Tokachi-oki earthquake in 339 Northern Japan (M8.1, 25 September 2003) and 340 San Simeon in Central California, M6.5, 22 De-341 cember 2003). The results, highly encouraging, are 342 described in Shebalin et al., in press. 343
- It seems natural to apply the RTP analysis to 344 earthquake precursors, expressed in other fields. 345 First positive results are obtained with precursors 346 gauging interaction between the ductile and brittle 347 layers of the crust (Aki, 2003; Jin et al., 2003; 348 Shebalin et al., in press). Other promising appli-349 cations include electromagnetic fields (Uyeda and 350 Park, 2002), fluid regime (Keilis-Borok, 1990; Ma 351 et al., 1990), GPS, InSAR, etc. 352
- We detect intermediate-term patterns only after a 353 chain has emerged so that its vicinity can be determined; this is too late to declare an intermediateterm alarm. Accordingly, our results concern only short-term prediction. 357
- "Pre-chain" precursors might emerge with a short 358 lead time too. 359
- There are no reasons not to explore RTP analysis 360 for prediction of different critical phenomena in hierarchical non-linear systems: other geological dis-362



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Fig. 5. Possible outcomes of prediction. Stars mark epicenters of strong earthquakes, targeted by prediction. A box to the right of the chain (dark gray) is the time–space covered by an alarm. A prediction is correct if a strong earthquake occurs within an alarm. Otherwise, this is a false alarm. Failure to predict is the case when a strong earthquake occurs outside of an alarm. Probabilistic component of prediction is represented by the rates of false alarms and failures to predict and the time–space covered by alarms (in % to total time–space considered).

asters; geotechnical, and even socio-economic disasters. Qualitatively similar approach is routinely
used in medicine, criminology, etc.

· However, accurate the short-term prediction would 366 be it will not render unnecessary the predictions 367 with a longer lead time. One can find in seismo-368 logical literature a reappearing mistake: that only 369 precise short-term (or even immediate) prediction is 370 371 practically useful. Actually, protection from earthquakes requires a hierarchy of preparedness mea-372 sures, from building codes, insurance, and issuing 373 bonds, to reinforcement of high risk objects, to red 374 375 alert. It takes different time, from decades, to years, 376 to seconds to undertake different measures. Accordingly, earthquake preparedness requires all stages 377 of prediction (Keilis-Borok, 2002; Molchan, 2003; 378 Kantorovich and Keilis-Borok, 1991). Such is the 379 380 case in preparedness to all disasters, war included.

381 4.5. Questions arising

• We considered only one short-term precursor—a chain of earthquakes—and one intermediate-term one—the pattern Σ . In subsequent applications (Shebalin et al., in press), all major types of intermediate-term seismicity patterns have been used with similar renormalization. The question arises which set of precursors provides the optimal prediction strategy, as defined for example in (Molchan, 2003; Zaliapin et al., 2003). 390

- It is not yet clear how to make the scaling of 391 RTP analysis self-adapting to the regional seismic 392 regime, e. g. to parameters of the Gutenberg–Richter 393 relation. 394
- Earthquake precursors emerge with the broader 395 range of the lead times than considered here. They 396 are divided, albeit fuzzily, into *long-term* (*tens of* 397 *years*) \Rightarrow *intermediate-term* (*years*) \Rightarrow *short-term* 398 (*months*) and \Rightarrow *immediate* (*days or less*). The 399 question arises how to apply RTP analysis to the 400 whole sequence or to its different parts. 401

Summing up, the RTP approach seems to open new402possibilities in the quest for the short-term prediction.403We hope that this study sets up a base for further devel-404opment of this approach in the intertwined problems405of earthquake prediction, fundamental understand-406ing of dynamics of the lithosphere, and non-linear407dynamics.408

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Gabrielov et al., 1999, Jordan, 2003, Zaliapin et al., 410 in press. 411

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