

1 **Reflection and transmission coefficients across a fracture**

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Abstract

15 The numerical modeling is used in many fields of science to explain the behavior of
16 some phenomena understandable. In geophysics, the simulations have application in wave
17 propagation, where the media can be characterized through their seismic responses. Due to
18 the computational cost, the models have to be as simple as possible, reducing the numerical
19 errors that these brings. In this work, we check the approximations (de qu ?) made by
20 Nakagawa and Schoenberg with the exact formulas, because of they simplify the proposed
21 models for media with fractures, characterizing its shear compliance, dry normal compliance
22 and membrane permeability. was the need to

23 The verification of the proposed models is done by comparing the reflection and trans-
24 mission coefficients obtained with the answer given by the model of the thin layer (para qu
25 cosa). In examples comparisons are made by varying the permeability and thickness of the
26 fracture; in the last example, three cases of interest were compared; in which the fracture
27 saturated with three different fluids, keeping the background with water.

28 Contrary to the paper mentioned above, the present paper is devoted to a justified
29 treatment of anisotropic elastic media with one, two, or three systems ...

30 FIGURAS EN EL PDF ADJUNTO, ESCRIBAN EL ABSTRACT, INTRODUCCION
31 Y CONCLUSIONES, Y DESPUES VEO

32 *SE VERIFICAN LAS APROXIMACIONES DE NAKAGAWA Y SCHOENBERG CON*
33 *LA FORMULA EXACTA. ESTO ES LO QUE SE HACE. HAY QUE DECIR PORQUE.*
34 *PORQUE SE VAN A IMPLEMENTAR EN MODELADO NUMERICO? OTRO MOTIVO?*
35 *HAY QUE SER CONVINCENTE. CUAL ES LA VENTAJA? EN PRINCIPIO SERIA NO*
36 *USAR CELDAS MUY PEQUENAS, ETC.*

37 **I. INTRODUCTION** Seismic wave propagation in fractured media is an active area of
38 research, with applications in many fields such as hydrocarbon geophysics exploration,
39 seismic monitoring of reservoir production and mining among others.

40 Modeling of fractures may be considered as a special case of the thin layer problem,
41 where the fracture is represented by a very thin layer with high permeability and
42 compliance.

43 There are relatively many works for a layer described by a single-phase (solid) case, e.
44 g., (41) and (2) consider the normal incidence case for a thin layer, (25) studied AVO
45 effects of a thin layer, while the effect of the thickness of a sedimentary layer has been
46 investigated by (17; 18).

47 (12) computes the scattering response of a lossy layer having orthorhombic symmetry
48 and embedded between two isotropic half-spaces, and (27) obtain the P-wave reflection
49 coefficient in isotropic lossless media as a function of the incidence angle.

50 Several theories have appeared in the literature to model fractures as boundary
51 conditions in the context of wave propagation phenomena.

52 The Linear-Slip Interface model (non-welded) for flat viscoelastic two-dimensional
53 fractures was proposed by Schoenberg, (40). This model imposes the continuity of the
54 stresses and discontinuity of the displacements across the fracture.

55 From laboratory experiences, (20; 21; Pyrak-Nolte et al. 1990) have validated the use

56 of this model. The experiments performed in (20; 21; Pyrak-Nolte et al. 1990) considered
57 the propagation of compressional and shear pulses in dry and wet fractured samples,
58 allowing to validate the seismic discontinuity displacement theory. Also, (28) have assumed
59 thin fractures as an elastic medium very soft in comparison with the frame.(agregar que
60 concluyo)

61 Concerning wave propagation in fractured fluid-saturated poroelastic media, we
62 mention the boundary conditions given by (?). and later by (?). In the latter reference,
63 several boundary conditions are developed and the corresponding reflection and
64 transmission coefficients are computed and analyzed. These boundary conditions first
65 consider the most general case in which stresses, velocities and fluid pressure may be
66 discontinuous across a fracture, and later several simplifying hypothesis allow to get other
67 forms of the boundary conditions, one of which reduces to that of (?).

68 In numerical simulations, to model fractures as very thin layers would require the use
69 of extremely fine computational meshes, and consequently employing boundary conditions
70 becomes a necessity.

71 In this paper we determine the frequency range in which the various boundary
72 conditions given in (?) are valid to represent fractures in numerical simulation of waves in
73 fractured poroelastic media.

74 For this purpose, we compare the reflection and transmission coefficients of waves

75 arriving to a plane fracture within a fluid-saturated porous medium represented either as a
76 thin layer or as boundary conditions.

77 The calculation of the Reflection and Transmission coefficients at a very thin
78 poroelastic separating two poroelastic half-spaces has been performed in (?). The results
79 were validated against limiting cases (elastic solid and inviscid fluids and zero layer
80 thickness) and the results predict all wave conversions, critical angles and polarity changes.

81 To our knowledge, the explicit calculation of the coefficients for poroelastic media has
82 not been addressed. Existing methods are restricted to normal incidence and/or are based
83 on numerical algorithms (1; 34; 35; 39). In general, these works are based on a
84 constitutive equation described by Biot's theory of poroelasticity (5; 6; 13; 14), which is
85 sufficiently general to model the desired characteristics of wave propagation, in particular,
86 the presence of the P waves (type-I and type-II compressional waves) and its effects on
87 interfaces (33? ? ?).

88 **There are also jobs where we consider two interfaces, as in (42), where the**
89 **system is formed by a fluid-saturated porous solid plate immersed in fluid, and**
90 **the results are compared with experimental data, and in (22), (?) and (19),**
91 **where ultrasonic measurements are compared with numerical data in**
92 **poroelastic slabs.**

93 We solve the scattering problem at all angles of incidence for a single layer embedded

94 between two half-spaces with dissimilar media, where the properties of the media are
95 described by Biot's theory of poroelasticity. The displacement fields are recast in terms of
96 potentials and the boundary conditions at the two interfaces impose continuity of the solid
97 and fluid displacements, normal and shear stresses and fluid pressure. The methodology is
98 analogous to that presented in (37), (36) and Carcione (12; 13), Section 6.4. The results
99 are verified for specific limiting cases with already published theoretical equations
100 (7; 13; 27; 32; 37).

101 The paper is organized as follows. Biot's theory is reviewed first. Then, we illustrate
102 the methodology and finally we present the examples. The final equations are verified with
103 limiting cases consisting of a single interface in poroelastic media and a layer, where the
104 media can be solids or fluids. The examples are relevant for applications in reflection
105 seismology.

106 Introduction

107 For examining the effect of pore fluids,

108 *GUITARRA. LA INTRODUCCION CONSISTE EN REVIEW, CLAIM Y LO QUE*
109 *SE HACE EN EL PAPER, 3 PARAGRAFOS*

110 II. BIOT'S THEORY

111 We consider a porous solid saturated by a viscous compressible fluid and assume that
112 the whole aggregate is isotropic. Let U and U_f be the averaged displacement vectors of the

113 solid and fluid parts of the medium, respectively. Then, W is defined as the averaged
 114 relative fluid displacement per unit volume of bulk material,

$$W = \phi(U_f - U), \quad (1)$$

115 where ϕ is the effective porosity.

116 Let ε_{ij} and σ_{ij} denote the strain tensors of the solid and the bulk material,
 117 respectively, and let P_f denote the fluid pressure. Following (5; 6), the stress-strain
 118 relations can be written as

$$\begin{aligned} \sigma_{ij} &= 2\mu\varepsilon_{ij}(U) + \delta_{ij}(\lambda_c\nabla\cdot U + D\nabla\cdot W), \quad i, j = 1, 2, 3, \\ P_f &= -D\nabla\cdot U - M\nabla\cdot W \end{aligned} \quad (2)$$

119 (13). Here, μ is the wet-rock shear modulus of the bulk material, considered to be equal to
 120 the shear modulus of the dry-rock. The grains are characterized by density ρ_s , bulk
 121 modulus K_s and shear modulus μ_s , while the fluid by ρ_f , K_f , and viscosity η . The grains
 122 are assumed to form an elastic porous matrix characterized by a porosity ϕ , permeability
 123 κ , bulk modulus K_m , and shear modulus μ_m . The Lamé constants of the saturated rock

124 are λ_c and μ . The constants λ_c , D and M in (2) can be written as (13)

$$\begin{aligned} \alpha &= 1 - \frac{K_m}{K_s}, \quad M = \left(\frac{\alpha - \phi}{K_s} + \frac{\phi}{K_f} \right)^{-1}, \quad D = \alpha M, \\ K_c &= K_m + \alpha^2 M, \quad \lambda_c = K_c - \frac{2}{3}\mu, \quad BK_u = D, \\ B &= \frac{1/K_m - 1/K_s}{1/K_m - 1/K_s + \phi(1/K_f - 1/K_s)}, \end{aligned} \quad (3)$$

125 where K_u is the undrained bulk modulus.

126 Next, let

$$\rho_b = (1 - \phi)\rho_s + \phi\rho_f \quad (4)$$

127 be the mass density of the bulk material. Also, let g and b denote the mass and viscous
128 coupling coefficients between the solid and fluid phases (3; 4):

$$g = \frac{S\rho_f}{\phi}, \quad b = \frac{\eta}{\kappa}, \quad S = \frac{1}{2} \left(1 + \frac{1}{\phi} \right), \quad (5)$$

129 where S is known as the structure factor. If g and b are functions of frequency, we have

$$\begin{aligned} b(\omega) &= \operatorname{Re} \left(\frac{\eta}{\kappa(\omega)} \right), \\ g(\omega) &= \frac{1}{\omega} \operatorname{Im} \left(\frac{\eta}{\kappa(\omega)} \right), \end{aligned} \quad (6)$$

130 being $\kappa(\omega)$ the dynamic permeability, a complex function defined in (23), and given by

$$\kappa(\omega) = \kappa_0 \left(\sqrt{1 + i \frac{4\omega}{n_j \omega_j}} + i \frac{\omega}{\omega_j} \right)^{-1}, \quad (7)$$

131 where κ_0 is the absolute permeability, n_j is a finite parameter determined by the pore
 132 geometry and ω_j is the viscous-boundary characteristic frequency given by
 133 $\omega_j = \eta\phi/(\kappa_0\rho_f S)$ (13; 23).

134 Then, assuming constant coefficients μ , λ_c , D , and M in (2), Biot's equations of
 135 motion can be stated as (5; 6; 13)

$$\begin{aligned}
 \nabla \cdot \sigma &= H_c \nabla (\nabla \cdot U) - \mu \nabla \times (\nabla \times U) + D \nabla (\nabla \cdot W) \\
 &= \rho_b \frac{\partial^2 U}{\partial t^2} + \rho_f \frac{\partial^2 W}{\partial t^2}, \\
 -\nabla P_f &= D \nabla (\nabla \cdot U) + M \nabla (\nabla \cdot W) \\
 &= \rho_f \frac{\partial^2 U}{\partial t^2} + g \frac{\partial^2 W}{\partial t^2} + b \frac{\partial W}{\partial t},
 \end{aligned} \tag{8}$$

136 where $H_c = \lambda_c + 2\mu$.

137 A plane-wave analysis shows that in this type of media two compressional waves
 138 (type-I and type-II waves) and one shear of S-wave can propagate (5).

139 III. REFLECTION AND TRANSMISSION COEFFICIENTS OF A SINGLE 140 LAYER

141 The fluid-saturated system consists of three media, Ω_n , $n = 1, 2, 3$ with different
 142 properties as shown in Figure 1. Let $z = 0$ be the boundary between Ω_1 and Ω_2 , and $z = h$
 143 the boundary between Ω_2 and Ω_3 , and consider a type-I compressional plane wave in Ω_1
 144 incident at $z = 0$ with an angle θ_{i1} with respect to the vertical z -axis. Following (38), we

145 represent the incident, reflected and transmitted waves using potentials.

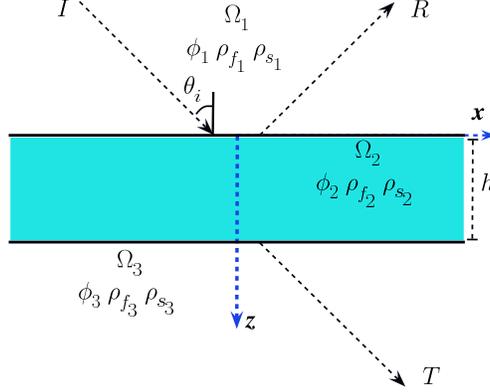


Figure 1: Geometry of the two half-spaces and the embedded layer.

146 For Ω_1 the potentials of the solid and relative fluid displacement are given by

$$\begin{aligned}\varphi_{i1} &= A_{i1} e^{i(\omega t - \mathbf{q}_{i1} \cdot \mathbf{x})}, \\ \psi_{i1} &= B_{i1} e^{i(\omega t - \mathbf{q}_{i1} \cdot \mathbf{x})},\end{aligned}\tag{9}$$

147 where

$$\mathbf{q}_{i1} = q_{i1} (\sin(\theta_{i1}), \cos(\theta_{i1})),$$

148 is the complex wave vector determining the polarization direction.

149 Let $\varphi_{rc}^{(1)}$, $\varphi_{rs}^{(1)}$, $\psi_{rc}^{(1)}$ and $\psi_{rs}^{(1)}$ be the compressional and shear potentials of the solid and

150 relative fluid displacement, for the reflected waves in Ω_1 . They are given by

$$\begin{aligned}
\varphi_{rc}^{(1)} &= A_{r1}^{(1)} e^{i(\omega t - \mathbf{q}_{r1}^{(1)} \cdot \mathbf{x})} + A_{r2}^{(1)} e^{i(\omega t - \mathbf{q}_{r2}^{(1)} \cdot \mathbf{x})}, \\
\varphi_{rs}^{(1)} &= A_{rs}^{(1)} e^{i(\omega t - \mathbf{q}_{rs}^{(1)} \cdot \mathbf{x})}, \\
\psi_{rc}^{(1)} &= B_{r1}^{(1)} e^{i(\omega t - \mathbf{q}_{r1}^{(1)} \cdot \mathbf{x})} + B_{r2}^{(1)} e^{i(\omega t - \mathbf{q}_{r2}^{(1)} \cdot \mathbf{x})}, \\
\psi_{rs}^{(1)} &= B_{rs}^{(1)} e^{i(\omega t - \mathbf{q}_{rs}^{(1)} \cdot \mathbf{x})},
\end{aligned} \tag{10}$$

151 where the subscript r indicates the reflected wave, c indicates compressional wave and s
152 shear wave, the super-index (1) refers to medium 1. Subscripts 1 and 2 indicate type-I and
153 type-II waves, respectively.

154 In Ω_2 , the potentials are

$$\begin{aligned}
\varphi_{tc}^{(2)} &= A_{t1}^{(2)} e^{i(\omega t - \mathbf{q}_{t1}^{(2)} \cdot \mathbf{x})} + A_{t2}^{(2)} e^{i(\omega t - \mathbf{q}_{t2}^{(2)} \cdot \mathbf{x})}, \\
\varphi_{ts}^{(2)} &= A_{ts}^{(2)} e^{i(\omega t - \mathbf{q}_{ts}^{(2)} \cdot \mathbf{x})}, \\
\psi_{tc}^{(2)} &= B_{t1}^{(2)} e^{i(\omega t - \mathbf{q}_{t1}^{(2)} \cdot \mathbf{x})} + B_{t2}^{(2)} e^{i(\omega t - \mathbf{q}_{t2}^{(2)} \cdot \mathbf{x})}, \\
\psi_{ts}^{(2)} &= B_{ts}^{(2)} e^{i(\omega t - \mathbf{q}_{ts}^{(2)} \cdot \mathbf{x})}, \\
\varphi_{rc}^{(2)} &= A_{r1}^{(2)} e^{i(\omega t - \mathbf{q}_{r1}^{(2)} \cdot \mathbf{x})} + A_{r2}^{(2)} e^{i(\omega t - \mathbf{q}_{r2}^{(2)} \cdot \mathbf{x})}, \\
\varphi_{rs}^{(2)} &= A_{rs}^{(2)} e^{i(\omega t - \mathbf{q}_{rs}^{(2)} \cdot \mathbf{x})}, \\
\psi_{rc}^{(2)} &= B_{r1}^{(2)} e^{i(\omega t - \mathbf{q}_{r1}^{(2)} \cdot \mathbf{x})} + B_{r2}^{(2)} e^{i(\omega t - \mathbf{q}_{r2}^{(2)} \cdot \mathbf{x})}, \\
\psi_{rs}^{(2)} &= B_{rs}^{(2)} e^{i(\omega t - \mathbf{q}_{rs}^{(2)} \cdot \mathbf{x})}
\end{aligned} \tag{11}$$

155 and the subscript t indicates the transmitted wave.

156 Finally, potentials in Ω_3 are expressed by

$$\begin{aligned}
\varphi_{tc}^{(3)} &= A_{t1}^{(3)} e^{i(\omega t - \mathbf{q}_{t1}^{(3)} \cdot \mathbf{x})} + A_{t2}^{(3)} e^{i(\omega t - \mathbf{q}_{t2}^{(3)} \cdot \mathbf{x})}, \\
\varphi_{ts}^{(3)} &= A_{ts}^{(3)} e^{i(\omega t - \mathbf{q}_{ts}^{(3)} \cdot \mathbf{x})}, \\
\psi_{tc}^{(3)} &= B_{t1}^{(3)} e^{i(\omega t - \mathbf{q}_{t1}^{(3)} \cdot \mathbf{x})} + B_{t2}^{(3)} e^{i(\omega t - \mathbf{q}_{t2}^{(3)} \cdot \mathbf{x})}, \\
\psi_{ts}^{(3)} &= B_{ts}^{(3)} e^{i(\omega t - \mathbf{q}_{ts}^{(3)} \cdot \mathbf{x})}.
\end{aligned} \tag{12}$$

157 In general, we determine $\mathbf{q}_{lj} = (\chi_{lj}, \beta_{lj}) = q_{lj}(\sin(\theta_{lj}), \cos(\theta_{lj}))$, $l = i, r, t$ and $j = 1, 2, s$
158 for each kind of wave.

159 The solid and relative fluid vectors $U^{(n)} = (U_x^{(n)}, U_z^{(n)})$ and $W^{(n)} = (W_x^{(n)}, W_z^{(n)})$ in
160 Ω_n , $n = 1, 2, 3$, are given by (37),

$$\begin{aligned}
U^{(1)} &= \nabla \varphi_{i1} + \nabla \varphi_{rc}^{(1)} + \left(-\frac{\partial \varphi_{rs}^{(1)}}{\partial z}, \frac{\partial \varphi_{rs}^{(1)}}{\partial x} \right), \\
&= U_{i1}^{(1)} + U_{r1}^{(1)} + U_{r2}^{(1)} + U_{rs}^{(1)}.
\end{aligned} \tag{13}$$

161

$$\begin{aligned}
W^{(1)} &= \nabla \psi_{i1} + \nabla \psi_{rc}^{(1)} + \left(-\frac{\partial \psi_{rs}^{(1)}}{\partial z}, \frac{\partial \psi_{rs}^{(1)}}{\partial x} \right), \\
&= W_{i1}^{(1)} + W_{r1}^{(1)} + W_{r2}^{(1)} + W_{rs}^{(1)}.
\end{aligned} \tag{14}$$

162

$$\begin{aligned}
U^{(2)} &= \nabla \varphi_{tc}^{(2)} + \left(-\frac{\partial \varphi_{ts}^{(2)}}{\partial z}, \frac{\partial \varphi_{ts}^{(2)}}{\partial x} \right) + \nabla \varphi_{rc}^{(2)} + \left(-\frac{\partial \varphi_{rs}^{(2)}}{\partial z}, \frac{\partial \varphi_{rs}^{(2)}}{\partial x} \right), \\
&= U_{t1}^{(2)} + U_{t2}^{(2)} + U_{ts}^{(2)} + U_{r1}^{(2)} + U_{r2}^{(2)} + U_{rs}^{(2)}.
\end{aligned} \tag{15}$$

163

$$\begin{aligned}
W^{(2)} &= \nabla\psi_{tc}^{(2)} + \left(-\frac{\partial\psi_{ts}^{(2)}}{\partial z}, \frac{\partial\psi_{ts}^{(2)}}{\partial x} \right) + \nabla\psi_{rc}^{(2)} + \left(-\frac{\partial\psi_{rs}^{(2)}}{\partial z}, \frac{\partial\psi_{rs}^{(2)}}{\partial x} \right), \\
&= W_{t1}^{(2)} + W_{t2}^{(2)} + W_{ts}^{(2)} + W_{r1}^{(2)} + W_{r2}^{(2)} + W_{rs}^{(2)}.
\end{aligned} \tag{16}$$

164

$$\begin{aligned}
U^{(3)} &= \nabla\varphi_{tc}^{(3)} + \left(-\frac{\partial\varphi_{ts}^{(3)}}{\partial z}, \frac{\partial\varphi_{ts}^{(3)}}{\partial x} \right), \\
&= U_{t1}^{(3)} + U_{t2}^{(3)} + U_{ts}^{(3)}.
\end{aligned} \tag{17}$$

165

$$\begin{aligned}
W^{(3)} &= \nabla\psi_{rc}^{(3)} + \left(-\frac{\partial\psi_{rs}^{(3)}}{\partial z}, \frac{\partial\psi_{rs}^{(3)}}{\partial x} \right), \\
&= W_{t1}^{(3)} + W_{t2}^{(3)} + W_{ts}^{(3)}.
\end{aligned} \tag{18}$$

166 Here $U_{lj}^{(n)}$ and $W_{lj}^{(n)}$, $l = i, r, t$, $j = 1, 2, s$, denote the type-I P wave, type-II P wave
167 and shear wave components of $U^{(n)}$ and $W^{(n)}$, respectively. The super-index (n) denotes
168 any variable associated with the medium Ω_n .

169 The boundary conditions at the interfaces located at $z = 0$ and $z = h$ impose
170 continuity of the solid and fluid displacements, continuity of the normal and shear stress
171 and continuity of the fluid pressure (38). Therefore, at $z = 0$ and $z = h$ we impose the

172 conditions

$$U_x^{(n)} = U_x^{(n+1)}, \quad (19)$$

$$U_z^{(n)} = U_z^{(n+1)}, \quad (20)$$

$$\sigma_{zz}^{(n)} = \sigma_{zz}^{(n+1)}, \quad (21)$$

$$\sigma_{xz}^{(n)} = \sigma_{xz}^{(n+1)}, \quad (22)$$

$$P_f^{(n)} = P_f^{(n+1)}, \quad (23)$$

$$W_z^{(n)} = W_z^{(n+1)}, \quad n = 1, 2. \quad (24)$$

173 The amplitude of the reflection and transmission coefficients $R_j^{(1)}$ and $T_j^{(3)}$, $j = 1, 2, s$,
 174 for the different waves are defined as the ratio of the solid-displacement amplitude of the
 175 corresponding wave and that of the incident wave (38), i.e,

$$R_j^{(1)} = \frac{A_{rj}^{(1)} q_{rj}^{(1)}}{A_{i1}^{(1)} q_{i1}^{(1)}}, \quad (25)$$

176 and

$$T_j^{(3)} = \frac{A_{tj}^{(3)} q_{tj}^{(3)}}{A_{i1}^{(1)} q_{i1}^{(1)}}. \quad (26)$$

177 Using equations (9)-(12) to obtain expressions for each of the pairs mentioned above
 178 and substituting them in (8) leads us to the following relationships between the amplitudes

179 of the solid and the relative amplitudes to the fluid (38):

$$\begin{aligned}
 B_{lj}^{(n)} &= \gamma_{lj}^{(n)} A_{lj}^{(n)}, \quad j = 1, 2, s, \quad l = r, t, \quad n = 1, 2, 3, \\
 B_{i1} &= \gamma_{i1} A_{i1},
 \end{aligned} \tag{27}$$

180 with

$$\begin{aligned}
 \gamma_{rj}^{(n)} &= \frac{\left[\rho_b^{(n)} \omega^2 - \left(q_{rj}^{(n)} \right)^2 H_c^{(n)} \right]}{\left[\left(q_{rj}^{(n)} \right)^2 D^{(n)} - \rho_f^{(n)} \omega^2 \right]} \quad j = 1, 2 \quad n = 1, 2, \\
 \gamma_{i1}^{(1)} &= \frac{\left[\rho_b^{(1)} \omega^2 - \left(q_{i1}^{(1)} \right)^2 H_c^{(1)} \right]}{\left[\left(q_{i1}^{(1)} \right)^2 D^{(1)} - \rho_f^{(1)} \omega^2 \right]}, \\
 \gamma_{tj}^{(n)} &= \frac{\left[\rho_b^{(n)} \omega^2 - \left(q_{tj}^{(n)} \right)^2 H_c^{(n)} \right]}{\left[\left(q_{tj}^{(n)} \right)^2 D^{(n)} - \rho_f^{(n)} \omega^2 \right]} \quad j = 1, 2 \quad n = 2, 3, \\
 \gamma_{rs}^{(n)} &= \frac{\mu^{(n)} \left(q_{rs}^{(n)} \right)^2 - \rho_b^{(n)} \omega^2}{\rho_f^{(n)} \omega^2} \quad n = 1, 2, \\
 \gamma_{ts}^{(n)} &= \frac{\mu^{(n)} \left(q_{ts}^{(n)} \right)^2 - \rho_b^{(n)} \omega^2}{\rho_f^{(n)} \omega^2} \quad n = 2, 3.
 \end{aligned}$$

181 The boundary conditions (19)-(24) require that the phase factors at the interfaces

182 $z = 0$ and $z = h$ are the same:

$$\begin{aligned}
 \chi_{i1} &= \chi_{r1}^{(1)} = \chi_{r2}^{(1)} = \chi_{rs}^{(1)} = \chi_{t1}^{(2)} = \chi_{t2}^{(2)} = \chi_{ts}^{(2)} = \chi_{r1}^{(2)} = \chi_{r2}^{(2)} = \chi_{rs}^{(2)} \\
 &= \chi_{t1}^{(3)} = \chi_{t2}^{(3)} = \chi_{ts}^{(3)} = \chi,
 \end{aligned} \tag{28}$$

183 which represents Snell's law and allows us to obtain the reflected and transmitted angles θ_{lj}
 184 for each type of wave as a function of the incidence angle θ_{i1} .

185 Application of the boundary conditions (19)-(24) and Snell's law (28) at $z = 0$ and
 186 $z = h$ give two systems of linear equations in the unknowns $A_{r1}, A_{r2}, A_{rs}, A_{t1}, A_{t2}$ and A_{ts}
 187 (see Appendix A). These two systems have coefficients depending on the wave numbers
 188 $q_{lj}^{(n)}, n = 1, 2, 3, l = i, r, t, j = 1, 2, s$.

189 Set

$$C_{lj}^{(n)} = A_{lj}^{(n)} / A_{i1}, \quad l = r, t, \quad j = 1, 2, s, \quad n = 1, 2, 3. \quad (29)$$

190 Using the matrix notation of Carcione (13), Section 6.4 to relate the fields at $z = 0$
 191 and $z = h$ we obtain

$$(\mathbf{A}_1 - \mathbf{B} * \mathbf{A}_3) \mathbf{r} = -\mathbf{i}_p, \quad (30)$$

192 where $\mathbf{r} = \left(C_{r1}^{(1)}, C_{r2}^{(1)}, C_{rs}^{(1)}, C_{t1}^{(3)}, C_{t2}^{(3)}, C_{ts}^{(3)} \right)^\top$,
 193 $\mathbf{i}_p = \left[-\chi, -\beta_{i1}^{(1)}, \zeta_{i1}^{(1)}, -2\mu^{(1)}\chi\beta_{i1}^{(1)}, \xi_{i1}^{(1)}, -\beta_{i1}^{(1)}\gamma_{i1}^{(1)} \right]^\top$ and $\mathbf{B} = \mathbf{T}(0) * (\mathbf{T}(h))^{-1}$ that acts as a
 194 boundary condition. The matrices of the system (30) are given in Appendix B.

195 The amplitude of the reflection and transmission coefficients for the different types of

196 waves are defined as

$$\begin{aligned} R_j^{(1)} &= C_{rj}^{(1)} \frac{q_{rj}^{(1)}}{q_{i1}^{(1)}}, \\ T_j^{(3)} &= C_{tj}^{(3)} \frac{q_{tj}^{(3)}}{q_{i1}^{(1)}}, \quad j = 1, 2, s. \end{aligned} \quad (31)$$

197 An incident S wave has the same scattering matrix as the P incident wave, but the
198 array \mathbf{i}_p in (30) is replaced by

$$\mathbf{i}_s = \left[\beta_{is}^{(1)}, -\chi, -\zeta_{is}^{(1)}, -\mu^{(1)} \left\{ \chi^2 - \left(\beta_{is}^{(1)} \right)^2 \right\}, 0, -\gamma_{is}^{(1)} \chi \right]^\top.$$

199 IV. SEISMIC BOUNDARY CONDITIONS ACROSS A FRACTURE

200 The model developed by (31) assumes that medium 1 and medium 3 have the same
201 properties and the thickness of medium 2 tends to zero ($h \rightarrow 0$). They obtain

$$\begin{bmatrix} \dot{U}_x^{(3)} - \dot{U}_x^{(1)} \\ \sigma_{zz}^{(3)} - \sigma_{zz}^{(1)} \\ (-P_f^{(3)}) - (-P_f^{(1)}) \\ \sigma_{xz}^{(3)} - \sigma_{xz}^{(1)} \\ \dot{U}_z^{(3)} - \dot{U}_z^{(1)} \\ \dot{W}_z^{(3)} - \dot{W}_z^{(1)} \end{bmatrix} = \frac{i\omega h}{2} \begin{bmatrix} \mathbf{0} & \tilde{\mathbf{Q}}_{XY} \\ \tilde{\mathbf{Q}}_{YX} & \mathbf{0} \end{bmatrix} \begin{bmatrix} \dot{U}_x^{(3)} + \dot{U}_x^{(1)} \\ \sigma_{zz}^{(3)} + \sigma_{zz}^{(1)} \\ (-P_f^{(3)}) + (-P_f^{(1)}) \\ \sigma_{xz}^{(3)} + \sigma_{xz}^{(1)} \\ \dot{U}_z^{(3)} + \dot{U}_z^{(1)} \\ \dot{W}_z^{(3)} + \dot{W}_z^{(1)} \end{bmatrix}, \quad (32)$$

202 where the dots over the displacement vector components indicate the time derivate. The
 203 matrices \mathbf{Q}_{XY} and \mathbf{Q}_{YX} are given by

$$\tilde{\mathbf{Q}}_{XY} = \begin{bmatrix} 1/G & \chi/\omega & 0 \\ \chi/\omega & \rho_b^{(2)} & \rho_f^{(2)} \cdot \Pi \\ 0 & \rho_f^{(2)} & \tilde{\rho}^{(2)} \cdot \Pi \end{bmatrix}, \quad (33)$$

204 and matrix $\tilde{\mathbf{Q}}_{YX}$ is gives explicitly as

$$\begin{aligned}
\tilde{\mathbf{Q}}_{YX}(1,1) &= -4\mu^{(2)} \left(\frac{\chi}{\omega}\right)^2 \left(1 - \frac{\mu^{(2)}}{H_c^{(2)}}\right) - \frac{\left(\rho_f^{(2)}\right)^2 - \rho_b^{(2)}\tilde{\rho}^{(2)}}{\tilde{\rho}^{(2)}} \\
&\quad - 2\mu^{(2)}\tilde{B}\tilde{\beta} \left(\frac{\chi}{\omega}\right)^2 \left(-\frac{\rho_f^{(2)}}{\tilde{\rho}^{(2)}} + \alpha^{(2)}\frac{2\mu^{(2)}}{H_c^{(2)}}\right) \cdot (1 - \Pi) \\
\tilde{\mathbf{Q}}_{YX}(1,2) &= \frac{\chi}{\omega} \left[\left(1 - \frac{2\mu^{(2)}}{H_c^{(2)}}\right) + \left(-\frac{\rho_f^{(2)}}{\tilde{\rho}^{(2)}} + \alpha^{(2)}\frac{2\mu^{(2)}}{H_c^{(2)}}\right) \tilde{B} \cdot (1 - \Pi) \right] \\
\tilde{\mathbf{Q}}_{YX}(1,3) &= \frac{\chi}{\omega} \left(-\frac{\rho_f^{(2)}}{\tilde{\rho}^{(2)}} + \alpha^{(2)}\frac{2\mu^{(2)}}{H_c^{(2)}}\right) \cdot \Pi \\
\tilde{\mathbf{Q}}_{YX}(2,1) &= \frac{\chi}{\omega} \left(1 - \frac{2\mu^{(2)}}{H_c^{(2)}} + 2\tilde{B}\tilde{\beta}\alpha^{(2)}\frac{\mu^{(2)}}{H_c^{(2)}} \cdot (1 - \Pi)\right) \\
\tilde{\mathbf{Q}}_{YX}(2,2) &= \frac{1}{H_c^{(2)}} - \alpha^{(2)}\tilde{B}\frac{1}{H_c^{(2)}} \cdot (1 - \Pi) \\
\tilde{\mathbf{Q}}_{YX}(2,3) &= -\alpha^{(2)}\frac{1}{H_c^{(2)}} \cdot \Pi \\
\tilde{\mathbf{Q}}_{YX}(3,1) &= \frac{\chi}{\omega} \left[-\frac{\rho_f^{(2)}}{\tilde{\rho}^{(2)}} + \alpha^{(2)}\frac{2\mu^{(2)}}{H_c^{(2)}} - 2\tilde{B}\tilde{\beta} \left((\alpha^{(2)})^2 \frac{\mu^{(2)}}{H_c^{(2)}} + \frac{\mu^{(2)}}{M^{(2)}} - \left(\frac{\chi}{\omega}\right)^2 \frac{\mu^{(2)}}{\tilde{\rho}^{(2)}} \right) \cdot (1 - \Pi) \right] \\
\tilde{\mathbf{Q}}_{YX}(3,2) &= -\alpha^{(2)}\frac{1}{H_c^{(2)}} + \left((\alpha^{(2)})^2 \frac{1}{H_c^{(2)}} + \frac{1}{M^{(2)}} - \left(\frac{\chi}{\omega}\right)^2 \frac{1}{\tilde{\rho}^{(2)}} \right) \tilde{B} \cdot (1 - \Pi) \\
\tilde{\mathbf{Q}}_{YX}(3,3) &= \left((\alpha^{(2)})^2 \frac{1}{H_c^{(2)}} + \frac{1}{M^{(2)}} - \left(\frac{\chi}{\omega}\right)^2 \frac{1}{\tilde{\rho}^{(2)}} \right) \cdot \Pi, \tag{34}
\end{aligned}$$

205

where

$$\begin{aligned}
\tilde{\rho} &= i \frac{\eta}{\omega \kappa(\omega)} \\
\frac{1}{\tilde{B}} &\equiv \alpha^{(2)} + \frac{H_c^{(2)}}{\alpha^{(2)} M^{(2)}} - \left(\frac{\chi}{\omega}\right)^2 \frac{H_c^{(2)}}{\alpha^{(2)} \tilde{\rho}^{(2)}}, \\
\tilde{\beta} &\equiv 1 - \frac{H_c^{(2)} \rho_f^{(2)}}{2 \alpha^{(2)} \mu^{(2)} \tilde{\rho}^{(2)}}, \\
\Pi &\equiv \frac{\tanh \epsilon}{\epsilon}, \quad \epsilon \equiv -\frac{i \beta_{t2}^{(2)} h}{2}.
\end{aligned} \tag{35}$$

206

The properties of medium 2 can be characterized by:

$$\eta_T \equiv \frac{h}{\mu^{(2)}} \quad (\text{shear compliance}), \tag{36}$$

$$\eta_{N_D} \equiv \frac{h}{H_c^{(2)}} \quad (\text{dry or drained normal compliance}), \tag{37}$$

$$\hat{\kappa}(\omega) \equiv \frac{\kappa^{(2)}(\omega)}{h} \quad (\text{membrane permeability}). \tag{38}$$

207

Simplifying equation (30), equation (52) of (31) is obtained:

$$\left\{ \begin{aligned}
&\dot{U}_x^{(3)} - \dot{U}_x^{(1)} = (i\omega) \eta_T \sigma_{xz}^{(1)} \\
&\dot{U}_z^{(3)} - \dot{U}_z^{(1)} = (i\omega) \eta_{N_D} \left[\left(1 - \alpha^{(2)} \tilde{B} (1 - \Pi)\right) \sigma_{zz}^{(1)} - \alpha^{(2)} \frac{-P_f^{(3)} + (-P_f^{(1)})}{2} \cdot \Pi \right] \\
&\dot{W}_z^{(3)} - \dot{W}_z^{(1)} = (i\omega) \alpha^{(2)} \eta_{N_D} \left[-\sigma_{zz}^{(1)} + \frac{1}{\tilde{B}} \frac{-P_f^{(3)} + (-P_f^{(1)})}{2} \right] \cdot \Pi \\
&\sigma_{xz}^{(3)} = \sigma_{xz}^{(1)} \\
&\sigma_{zz}^{(3)} = \sigma_{zz}^{(1)} \\
&-P_f^{(3)} - (-P_f^{(1)}) = \frac{\eta_f^{(2)}}{\hat{\kappa}(\omega)} \frac{\dot{W}_z^{(3)} + \dot{W}_z^{(1)}}{2} \cdot \Pi
\end{aligned} \right. \tag{39}$$

208

where $\tilde{\beta} \approx 1$ and

$$\begin{aligned}
\tilde{B} &= \alpha^{(2)} \frac{M^{(2)}}{H_u^{(2)}}, \\
H_u^{(2)} &= K_u^{(2)} + \frac{4\mu^{(2)}}{3}, \\
\epsilon &= \frac{1-i}{2} \sqrt{\omega \frac{\alpha^{(2)} \eta_f \eta_{ND}}{2 * \tilde{B} \hat{\kappa}_0}}, \quad \hat{\kappa}_0 = \kappa_0/h.
\end{aligned} \tag{40}$$

209

When the permeability tends to infinity, equation (53) of (31) is obtained.

$$\left\{ \begin{array}{l}
\dot{U}_x^{(3)} - \dot{U}_x^{(1)} = (i\omega) \eta_T \sigma_{xz}^{(1)} \\
\dot{U}_z^{(3)} - \dot{U}_z^{(1)} = (i\omega) \eta_{ND} \left[\sigma_{zz}^{(1)} - \alpha^{(2)} \left(-P_f^{(1)} \right) \right] \\
\dot{W}_z^{(3)} - \dot{W}_z^{(1)} = (i\omega) \alpha^{(2)} \eta_{ND} \left[-\sigma_{zz}^{(1)} + \frac{1}{B} \left(-P_f^{(1)} \right) \right] \\
\sigma_{xz}^{(3)} = \sigma_{xz}^{(1)} \\
\sigma_{zz}^{(3)} = \sigma_{zz}^{(1)} \\
-P_f^{(3)} = -P_f^{(1)}
\end{array} \right. \tag{41}$$

210 V. EXAMPLES

211

We model the fracture as a thin layer whose reflection coefficient is given by equation

212

(30) assuming h much smaller than the signal wavelength (thin-layer model or TL model).

213

Let us define the fracture-thickness/wavelength ratio, \mathcal{R} , where the wavelength is that of

214

the background medium. We consider several cases of interest in reservoir geophysics.

215

Media properties and fracture properties are shown in Table 1 and fluid properties are

216

shown in Table 2. The following cases are taken into account:

217 **Case 1:** Comparison between the TL model and equation (30) of (31).

218 **Case 2:** Comparison between the TL model and equation (52) of (31).

219 **Case 3:** Comparison between the TL model and equation (53) of (31).

220 **Case 4:** Calculation of reflection and transmission coefficients for three types of fluid
 221 in the fracture, with water in the background medium, using the equation (52) of
 222 (31).

223 **A. Case 1**

224 We compute the reflection and transmission coefficients for compressional plane waves
 225 propagating through a fracture, where the fluid is water everywhere. Considering the
 226 notation in Figure 1, medium 2 is the fracture.

227 Figure 2 shows the magnitude of the reflection and transmission coefficients for the
 228 TL model and equation (30) of (31), where $\mathcal{R} = 3.1 \times 10^{-4}$. When the permeability is very
 229 small, there are differences between the two models. In Figure 3, $\mathcal{R} = 3.1 \times 10^{-6}$ and here
 230 the coefficients obtained with the two models are the same except at very high frequencies.

231 **B. Case 2**

232 In this case, the TL model is compared with equation (52) of (31). The permeability
 233 is $\kappa_0 = 10^{-4} D$, the incident wave is a type I P-wave and the thickness of the fracture varies

234 from $h = 0.001$ m to $h = 0.00001$ m. When $\mathcal{R} = 3.1 \times 10^{-5}$, at approximately 100 Hz,
 235 there are differences between the two models, see Figure 4. But these differences disappear
 236 when $\mathcal{R} = 3.1 \times 10^{-7}$.

237 C. Case 3

238 In this case, the permeability is infinite ($\kappa_0 \rightarrow \infty$). When $h = 0.001$ m and the
 239 frequency is 1000 Hz, there are differences in the absolute value of the transmission
 240 coefficient obtained with the TL model and equation (53) of (31) when $\mathcal{R} = 3.1 \times 10^{-4}$. If
 241 $h = 0.00001$ m there are no differences between the two models. In this case, it is $\mathcal{R} = 3.1$
 242 $\times 10^{-6}$ for a frequency of 1000 Hz.

243 D. Case 4

244 After comparing the present model with that of (31), we proceed to analyze three
 245 different cases of interest in reservoir geophysics. The medium is saturated with water and
 246 the fracture contains three different fluids, water, oil and gas. The fluid properties are
 247 shown in Table 2. The coefficients were calculated for a frequency of 50 Hz and fracture
 248 thickness $h = 0.001$ m, which gives $\mathcal{R} = 1.6 \times 10^{-5}$

249 The reflection and transmission coefficients are calculated with equation (52) of (31)
 250 and the results are shown in Figure 6. It can be seen that when the bulk modulus of the
 251 fluid in the fracture differs from that of the background, the coefficients vary more. When

252 the fluid density in the fracture is much less than that of the background, the two peaks
253 appearing in the reflection coefficient of the type I wave (R_{PI}) are closer to each other. Also
254 displayed is a peak at 50 degrees in the coefficients of the wave type II (R_{PII} and T_{PII})
255 when there is gas in the fracture.

256 VI. CONCLUSIONS

257 In all cases shown, when the thickness is too small the fracture ratio / wavelength, the
258 results obtained with (31) approximations are similar to the results obtained with the fine
259 layer model, especially for low frequencies . A variable that affects the results very
260 significantly, is the permeability of the fracture; to cases with low permeability fractures,
261 the thickness must be very small, otherwise the approximations and the fine layer model
262 differ, especially at high frequencies. This shows that in certain cases the approximations
263 do not fit the exact model.

264 In the latter case it is observed that when there is gas in the fracture, the coefficient
265 of reflection of PI-wave, is very different from the other two cases (water or oil in the
266 fracture) and has two lobes between 40 and 60 degrees. Something similar happens with
267 the coefficients of reflection and transmission of shear wave, which appears lobe near 50
268 degrees only for the case of gas in the fracture. This could be used as an indicator of gas in
269 fractures.

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274 APPENDIX A: LINEAR SYSTEMS

275 Here, we report the linear equations for the unknown amplitude of the reflected and
 276 transmitted waves. First, application of the boundary conditions (19)-(24) at $z = 0$ yields
 277 the linear system

$$\begin{aligned}
 -\chi A_{i1}^{(1)} - \chi A_{r1}^{(1)} - \chi A_{r2}^{(1)} + \beta_{rs}^{(1)} A_{rs}^{(1)} = -\chi A_{t1}^{(2)} - \chi A_{t2}^{(2)} + \beta_{ts}^{(2)} A_{ts}^{(2)} - \chi A_{r1}^{(2)} \\
 -\chi A_{r2}^{(2)} + \beta_{rs}^{(2)} A_{rs}^{(2)}. \quad (\text{A-1})
 \end{aligned}$$

278

$$\begin{aligned}
 -\beta_{i1}^{(1)} A_{i1}^{(1)} - \beta_{r1}^{(1)} A_{r1}^{(1)} - \beta_{r2}^{(1)} A_{r2}^{(1)} - \chi A_{rs}^{(1)} = -\beta_{t1}^{(2)} A_{t1}^{(2)} - \beta_{t2}^{(2)} A_{t2}^{(2)} - \chi A_{ts}^{(2)} \\
 -\beta_{r1}^{(2)} A_{r1}^{(2)} - \beta_{r2}^{(2)} A_{r2}^{(2)} - \chi A_{rs}^{(2)}. \quad (\text{A-2})
 \end{aligned}$$

279

$$\begin{aligned}
 A_{i1}^{(1)} \zeta_{i1}^{(1)} + A_{r1}^{(1)} \zeta_{r1}^{(1)} + A_{r2}^{(1)} \zeta_{r2}^{(1)} - A_{rs}^{(1)} \zeta_{rs}^{(1)} = A_{t1}^{(2)} \zeta_{t1}^{(2)} + A_{t2}^{(2)} \zeta_{t2}^{(2)} - A_{ts}^{(2)} \zeta_{ts}^{(2)} \\
 + A_{r1}^{(2)} \zeta_{r1}^{(2)} + A_{r2}^{(2)} \zeta_{r2}^{(2)} - A_{rs}^{(2)} \zeta_{rs}^{(2)}. \quad (\text{A-3})
 \end{aligned}$$

280

$$\begin{aligned}
 -2\mu^{(1)} A_{i1} \chi \beta_{i1}^{(1)} - 2\mu^{(1)} A_{r1} \chi \beta_{r1}^{(1)} - 2\mu^{(1)} A_{r2} \chi \beta_{r2}^{(1)} - \mu^{(1)} A_{rs} \left[\chi^2 - (\beta_{rs}^{(1)})^2 \right] = \\
 -2\mu^{(2)} A_{t1} \chi \beta_{t1}^{(2)} - 2\mu^{(2)} A_{t2} \chi \beta_{t2}^{(2)} - \mu^{(2)} A_{ts} \left[\chi^2 - (\beta_{ts}^{(2)})^2 \right] - 2\mu^{(2)} A_{r1} \chi \beta_{r1}^{(2)} \\
 -2\mu^{(2)} A_{r2} \chi \beta_{r2}^{(2)} - \mu^{(2)} A_{rs} \left[\chi^2 - (\beta_{rs}^{(2)})^2 \right]. \quad (\text{A-4})
 \end{aligned}$$

281

$$A_{i1}^{(1)} \zeta_{i1}^{(1)} + A_{r1}^{(1)} \zeta_{r1}^{(1)} + A_{r2}^{(1)} \zeta_{r2}^{(1)} = A_{t1}^{(2)} \zeta_{t1}^{(2)} + A_{t2}^{(2)} \zeta_{t2}^{(2)} + A_{r1}^{(2)} \zeta_{r1}^{(2)} + A_{r2}^{(2)} \zeta_{r2}^{(2)}. \quad (\text{A-5})$$

282

$$\begin{aligned} -\beta_{i1}^{(1)} \gamma_2^{(1)} A_{i1}^{(1)} - \beta_{r1}^{(1)} \gamma_1^{(1)} A_{r1}^{(1)} - \beta_{r2}^{(1)} \gamma_2^{(1)} A_{r2}^{(1)} - \chi \gamma_{rs}^{(1)} A_{rs}^{(1)} &= -\beta_{t1}^{(2)} \gamma_1^{(2)} A_{t1}^{(2)} \\ -\beta_{t2}^{(2)} \gamma_2^{(2)} A_{t2}^{(2)} - \chi \gamma_{ts}^{(2)} A_{ts}^{(2)} - \beta_{r1}^{(2)} \gamma_1^{(2)} A_{r1}^{(2)} - \beta_{r2}^{(2)} \gamma_2^{(2)} A_{r2}^{(2)} - \chi \gamma_{rs}^{(2)} A_{rs}^{(2)}. \end{aligned} \quad (\text{A-6})$$

283

Similarly, at $z = h$ we obtain

$$\begin{aligned} -\chi A_{t1}^{(3)} e^{-i\beta_{t1}^{(3)} h} - \chi A_{t2}^{(3)} e^{-i\beta_{t2}^{(3)} h} + \beta_{ts}^{(3)} A_{ts}^{(3)} e^{-i\beta_{ts}^{(3)} h} &= -\chi A_{t1}^{(2)} e^{-i\beta_{t1}^{(2)} h} \\ -\chi A_{t2}^{(2)} e^{-i\beta_{t2}^{(2)} h} + \beta_{ts}^{(2)} A_{ts}^{(2)} e^{-i\beta_{ts}^{(2)} h} - \chi A_{r1}^{(2)} e^{-i\beta_{r1}^{(2)} h} - \chi A_{r2}^{(2)} e^{-i\beta_{r2}^{(2)} h} \\ &\quad + \beta_{rs}^{(2)} A_{rs}^{(2)} e^{-i\beta_{rs}^{(2)} h}. \end{aligned} \quad (\text{A-7})$$

284

$$\begin{aligned} -\beta_{t1}^{(3)} A_{t1}^{(3)} e^{-i\beta_{t1}^{(3)} h} - \beta_{t2}^{(3)} A_{t2}^{(3)} e^{-i\beta_{t2}^{(3)} h} - \chi A_{ts}^{(3)} e^{-i\beta_{ts}^{(3)} h} &= -\beta_{t1}^{(2)} A_{t1}^{(2)} e^{-i\beta_{t1}^{(2)} h} \\ -\beta_{t2}^{(2)} A_{t2}^{(2)} e^{-i\beta_{t2}^{(2)} h} - \chi A_{ts}^{(2)} e^{-i\beta_{ts}^{(2)} h} - \beta_{r1}^{(2)} A_{r1}^{(2)} e^{-i\beta_{r1}^{(2)} h} - \beta_{r2}^{(2)} A_{r2}^{(2)} e^{-i\beta_{r2}^{(2)} h} \\ &\quad - \chi A_{rs}^{(2)} e^{-i\beta_{rs}^{(2)} h}. \end{aligned} \quad (\text{A-8})$$

285

$$\begin{aligned} A_{t1}^{(3)} \zeta_{t1}^{(3)} e^{-i\beta_{t1}^{(3)} h} + A_{t2}^{(3)} \zeta_{t2}^{(3)} e^{-i\beta_{t2}^{(3)} h} - A_{ts}^{(3)} \zeta_{ts}^{(3)} e^{-i\beta_{ts}^{(3)} h} &= A_{t1}^{(2)} \zeta_{t1}^{(2)} e^{-i\beta_{t1}^{(2)} h} \\ + A_{t2}^{(2)} \zeta_{t2}^{(2)} e^{-i\beta_{t2}^{(2)} h} - A_{ts}^{(2)} \zeta_{ts}^{(2)} e^{-i\beta_{ts}^{(2)} h} + A_{r1}^{(2)} \zeta_{r1}^{(2)} e^{-i\beta_{r1}^{(2)} h} + A_{r2}^{(2)} \zeta_{r2}^{(2)} e^{-i\beta_{r2}^{(2)} h} \\ &\quad - A_{rs}^{(2)} \zeta_{rs}^{(2)} e^{-i\beta_{rs}^{(2)} h}. \end{aligned} \quad (\text{A-9})$$

286

$$\begin{aligned}
& -2\mu^{(3)} A_{t1} \chi \beta_{t1}^{(3)} e^{-i\beta_{t1}^{(3)} h} - 2\mu^{(3)} A_{t2} \chi \beta_{t2}^{(3)} e^{-i\beta_{t2}^{(3)} h} - \mu^{(3)} A_{ts} \left[\chi^2 - \left(\beta_{ts}^{(3)} \right)^2 \right] e^{-i\beta_{ts}^{(3)} h} = \\
& -2\mu^{(2)} A_{t1}^{(2)} \chi \beta_{t1}^{(2)} e^{-i\beta_{t1}^{(2)} h} - 2\mu^{(2)} A_{t2}^{(2)} \chi \beta_{t2}^{(2)} e^{-i\beta_{t2}^{(2)} h} - \mu^{(2)} A_{ts}^{(2)} \left[\chi^2 - \left(\beta_{ts}^{(2)} \right)^2 \right] e^{-i\beta_{ts}^{(2)} h} \\
& -2\mu^{(2)} A_{r1}^{(2)} \chi \beta_{r1}^{(2)} e^{-i\beta_{r1}^{(2)} h} - 2\mu^{(2)} A_{r2}^{(2)} \chi \beta_{r2}^{(2)} e^{-i\beta_{r2}^{(2)} h} - \mu^{(2)} A_{rs}^{(2)} \left[\chi^2 - \left(\beta_{rs}^{(2)} \right)^2 \right] e^{-i\beta_{rs}^{(2)} h}. \quad (\text{A-10})
\end{aligned}$$

287

$$\begin{aligned}
A_{t1}^{(3)} \xi_{t1}^{(3)} e^{-i\beta_{t1}^{(3)} h} + A_{t2}^{(3)} \xi_{t2}^{(3)} e^{-i\beta_{t2}^{(3)} h} &= A_{t1}^{(2)} \xi_{t1}^{(2)} e^{-i\beta_{t1}^{(2)} h} + A_{t2}^{(2)} \xi_{t2}^{(2)} e^{-i\beta_{t2}^{(2)} h} \\
&+ A_{r1}^{(2)} \xi_{r1}^{(2)} e^{-i\beta_{r1}^{(2)} h} + A_{r2}^{(2)} \xi_{r2}^{(2)} e^{-i\beta_{r2}^{(2)} h}. \quad (\text{A-11})
\end{aligned}$$

288

$$\begin{aligned}
& -\beta_{t1}^{(3)} \gamma_1^{(3)} A_{t1}^{(3)} e^{-i\beta_{t1}^{(3)} h} - \beta_{t2}^{(3)} \gamma_2^{(3)} A_{t2}^{(3)} e^{-i\beta_{t2}^{(3)} h} - \chi \gamma_{ts}^{(3)} A_{ts}^{(3)} e^{-i\beta_{ts}^{(3)} h} = \\
& -\beta_{t1}^{(2)} \gamma_1^{(2)} A_{t1}^{(2)} e^{-i\beta_{t1}^{(2)} h} - \beta_{t2}^{(2)} \gamma_2^{(2)} A_{t2}^{(2)} e^{-i\beta_{t2}^{(2)} h} - \chi \gamma_{ts}^{(2)} A_{ts}^{(2)} e^{-i\beta_{ts}^{(2)} h} \\
& -\beta_{r1}^{(2)} \gamma_1^{(2)} A_{r1}^{(2)} e^{-i\beta_{r1}^{(2)} h} - \beta_{r2}^{(2)} \gamma_2^{(2)} A_{r2}^{(2)} e^{-i\beta_{r2}^{(2)} h} - \chi \gamma_{rs}^{(2)} A_{rs}^{(2)} e^{-i\beta_{rs}^{(2)} h}. \quad (\text{A-12})
\end{aligned}$$

289

The coefficients of the systems (A-1)-(A-6) and (A-7)-(A-12) are given by

$$\begin{aligned}
\beta_{rj}^{(n)} &= -\sqrt{\left(q_{rj}^{(n)} \right)^2 - \chi^2}, \quad j = 1, 2, s \quad n = 1, 2, \\
\beta_{tj}^{(n)} &= \sqrt{\left(q_{tj}^{(n)} \right)^2 - \chi^2}, \quad j = 1, 2, s \quad n = 2, 3, \\
\zeta_{lj}^{(n)} &= -\left(q_{lj}^{(n)} \right)^2 \left(H_c^{(n)} + D^{(n)} \gamma_{lj}^{(n)} \right) + 2\mu^{(n)} \chi^2, \quad j = 1, 2 \quad n = 1, 2, 3 \quad l = i, r, t, \\
\zeta_{ls}^{(n)} &= 2\mu^{(n)} \chi \beta_{ls}^{(n)}, \quad n = 1, 2, 3 \quad l = r, t, \\
\xi_{lj}^{(n)} &= \left(D^{(n)} + M^{(n)} \gamma_{lj}^{(n)} \right) \left(q_{lj}^{(n)} \right)^2, \quad j = 1, 2 \quad n = 1, 2, 3 \quad l = i, r, t.
\end{aligned} \quad (\text{A-13})$$

290 Now, using equations (3) and (5), we obtain the wave numbers $q_{lj}^{(n)}$, $n = 1, 2, 3$,

291 $l = i, r, t$, and $j = 1, 2, s$:

$$\begin{aligned}
 q_{r1}^{(n)} &= \sqrt{\frac{-F^{(n)} - \sqrt{(F^{(n)})^2 - 4G^{(n)}K^{(n)}}}{2G^{(n)}}}, \quad n = 1, 2, 3, \\
 q_{r2}^{(n)} &= \sqrt{\frac{-F^{(n)} + \sqrt{(F^{(n)})^2 - 4G^{(n)}K^{(n)}}}{2G^{(n)}}}, \quad n = 1, 2, 3, \\
 q_{rs}^{(n)} &= \sqrt{N^{(n)} - iV^{(n)}}, \quad n = 1, 2, 3, \\
 q_{r1}^{(1)} &= q_{i1}^{(1)}, \\
 q_{rj}^{(n)} &= q_{tj}^{(n)}, \quad j = 1, 2, s \quad n = 1, 2, 3.
 \end{aligned} \tag{A-14}$$

292 Here

$$\begin{aligned}
 G^{(n)} &= M^{(n)}H_c^{(n)} - (D^{(n)})^2, \quad n = 1, 2, 3, \\
 F^{(n)} &= \omega^2 \left[2\rho_f^{(n)}D^{(n)} - H_c^{(n)}g^{(n)} - \rho_b^{(n)}M^{(n)} \right] + iH_c^{(n)}b^{(n)}\omega, \quad n = 1, 2, 3, \\
 K^{(n)} &= \omega^4 \left[\rho_b^{(n)}g^{(n)} - (\rho_f^{(n)})^2 \right] - i\rho_b^{(n)}b^{(n)}\omega^3, \quad n = 1, 2, 3, \\
 N^{(n)} &= \frac{\omega^2}{\mu^{(n)}} \left[\rho_b^{(n)} - \frac{(\rho_f^{(n)})^2\omega^2g^{(n)}}{(g^{(n)})^2\omega^2 + (b^{(n)})^2} \right], \quad n = 1, 2, 3, \\
 V^{(n)} &= (\rho_f^{(n)})^2 \frac{\omega^3b^{(n)}}{\mu^{(n)}((g^{(n)})^2\omega^2 + (b^{(n)})^2)}, \quad n = 1, 2, 3.
 \end{aligned}$$

293 **APPENDIX B: FINAL SYSTEM OF EQUATIONS**

294

Each of the matrices in the system (30) are defined by

$$\mathbf{A}_1 = \begin{pmatrix} -\chi & -\chi & \beta_{rs}^{(1)} & 0 & 0 & 0 \\ -\beta_{r1}^{(1)} & -\beta_{r2}^{(1)} & -\chi & 0 & 0 & 0 \\ \zeta_{r1}^{(1)} & \zeta_{r2}^{(1)} & -\zeta_{rs}^{(1)} & 0 & 0 & 0 \\ -2\mu^{(1)}\chi\beta_{r1}^{(1)} & -2\mu^{(1)}\chi\beta_{r2}^{(1)} & -\mu^{(1)}\left[\chi^2 - (\beta_{rs}^{(1)})^2\right] & 0 & 0 & 0 \\ \xi_{r1}^{(1)} & \xi_{r2}^{(1)} & 0 & 0 & 0 & 0 \\ -\gamma_{r1}^{(1)}\beta_{r1}^{(1)} & -\gamma_{r2}^{(1)}\beta_{r2}^{(1)} & -\gamma_{rs}^{(1)}\chi & 0 & 0 & 0 \end{pmatrix}, \quad (\text{B-1})$$

295 and

$$\mathbf{A}_3 = \begin{pmatrix} 0 & 0 & 0 \\ 0 & 0 & 0 \\ 0 & 0 & 0 \\ 0 & 0 & 0 \\ 0 & 0 & 0 \\ 0 & 0 & 0 \end{pmatrix}$$

$$\begin{pmatrix}
-\chi e^{-i\beta_{t1}^{(3)}h} & -\chi e^{-i\beta_{t2}^{(3)}h} & \beta_{ts}^{(3)} e^{-i\beta_{ts}^{(3)}h} \\
-\beta_{t1}^{(3)} e^{-i\beta_{t1}^{(3)}h} & -\beta_{t2}^{(3)} e^{-i\beta_{t2}^{(3)}h} & -\chi e^{-i\beta_{ts}^{(3)}h} \\
\zeta_{t1}^{(3)} e^{-i\beta_{t1}^{(3)}h} & \zeta_{t2}^{(3)} e^{-i\beta_{t2}^{(3)}h} & -\zeta_{ts}^{(3)} e^{-i\beta_{ts}^{(3)}h} \\
-2\mu^{(3)} \chi \beta_{t1}^{(3)} e^{-i\beta_{t1}^{(3)}h} & -2\mu^{(3)} \chi \beta_{t2}^{(3)} e^{-i\beta_{t2}^{(3)}h} & -\mu^{(3)} \left[\chi^2 - \left(\beta_{ts}^{(3)} \right)^2 \right] e^{-i\beta_{ts}^{(3)}h} \\
\xi_{t1}^{(3)} e^{-i\beta_{t1}^{(3)}h} & \xi_{t2}^{(3)} e^{-i\beta_{t2}^{(3)}h} & 0 \\
-\gamma_{t1}^{(3)} \beta_{t1}^{(3)} e^{-i\beta_{t1}^{(3)}h} & -\gamma_{t2}^{(3)} \beta_{t2}^{(3)} e^{-i\beta_{t2}^{(3)}h} & -\gamma_{ts}^{(3)} \chi e^{-i\beta_{ts}^{(3)}h}
\end{pmatrix}. \quad (\text{B-2})$$

Finally, $\mathbf{B} = \mathbf{T}(0) * (\mathbf{T}(h))^{-1}$, and $T(z) = \mathbf{S}_1 * \mathbf{S}_2(z)$, being:

$$\mathbf{S}_1 = \begin{pmatrix}
-\chi & -\chi & \beta_{ts}^{(2)} \\
-\beta_{t1}^{(2)} & -\beta_{t2}^{(2)} & -\chi \\
\zeta_{t1}^{(2)} & \zeta_{t2}^{(2)} & -\zeta_{ts}^{(2)} \\
-2\mu^{(2)} \chi \beta_{t1}^{(2)} & -2\mu^{(2)} \chi \beta_{t2}^{(2)} & -\mu^{(2)} \left[\chi^2 - \left(\beta_{ts}^{(2)} \right)^2 \right] \\
\xi_{t1}^{(2)} & \xi_{t2}^{(2)} & 0 \\
-\gamma_{t1}^{(2)} \beta_{t1}^{(2)} & -\gamma_{t2}^{(2)} \beta_{t2}^{(2)} & -\gamma_{ts}^{(2)} \chi
\end{pmatrix}$$

298

$$\left. \begin{array}{ccc}
 -\chi & -\chi & \beta_{rs}^{(2)} \\
 -\beta_{r1}^{(2)} & -\beta_{r2}^{(2)} & -\chi \\
 \zeta_{r1}^{(2)} & \zeta_{r2}^{(2)} & -\zeta_{rs}^{(2)} \\
 -2\mu^{(2)}\chi\beta_{r1}^{(2)} & -2\mu^{(2)}\chi\beta_{r2}^{(2)} & -\mu^{(2)}\left[\chi^2 - \left(\beta_{rs}^{(2)}\right)^2\right] \\
 \xi_{r1}^{(2)} & \xi_{r2}^{(2)} & 0 \\
 -\gamma_{r1}^{(2)}\beta_{r1}^{(2)} & -\gamma_{r2}^{(2)}\beta_{r2}^{(2)} & -\gamma_{rs}^{(2)}\chi
 \end{array} \right), \quad (\text{B-3})$$

299 and

$$\mathbf{S}_2(z) = \left(\begin{array}{cccccc}
 e^{-i\beta_{t1}^{(2)}z} & 0 & 0 & 0 & 0 & 0 \\
 0 & e^{-i\beta_{t2}^{(2)}z} & 0 & 0 & 0 & 0 \\
 0 & 0 & e^{-i\beta_{ts}^{(2)}z} & 0 & 0 & 0 \\
 0 & 0 & 0 & e^{-i\beta_{r1}^{(2)}z} & 0 & 0 \\
 0 & 0 & 0 & 0 & e^{-i\beta_{r2}^{(2)}z} & 0 \\
 0 & 0 & 0 & 0 & 0 & e^{-i\beta_{rs}^{(2)}z}
 \end{array} \right). \quad (\text{B-4})$$

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Table 1: Baseline material properties used for the numerical examples are shown.

Properties	Matrix	Fracture
Porosity	0.15	0.5
Solid density (g/m ³)	2.7	2.7
Solid bulk modulus (GPa)	36	36
Frame bulk modulus (GPa)	9	0.0556
Frame shear modulus (GPa)	7	0.0333
Permeability (D)	0.1	100 (case 1, 4)
		0.0001 (case 1)
		0.001 (case 2)
		∞ (case 3)
Tortuosity	3	1

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Table 2: Saturant fluids.

Properties	Gas	Water	Oil
Density (g/m ³)	0.1398	1	0.7
Fluid viscosity (Pa s)	0.000022	0.001	0.004
Fluid bulk modulus (GPa)	0.05543	2.25	0.57

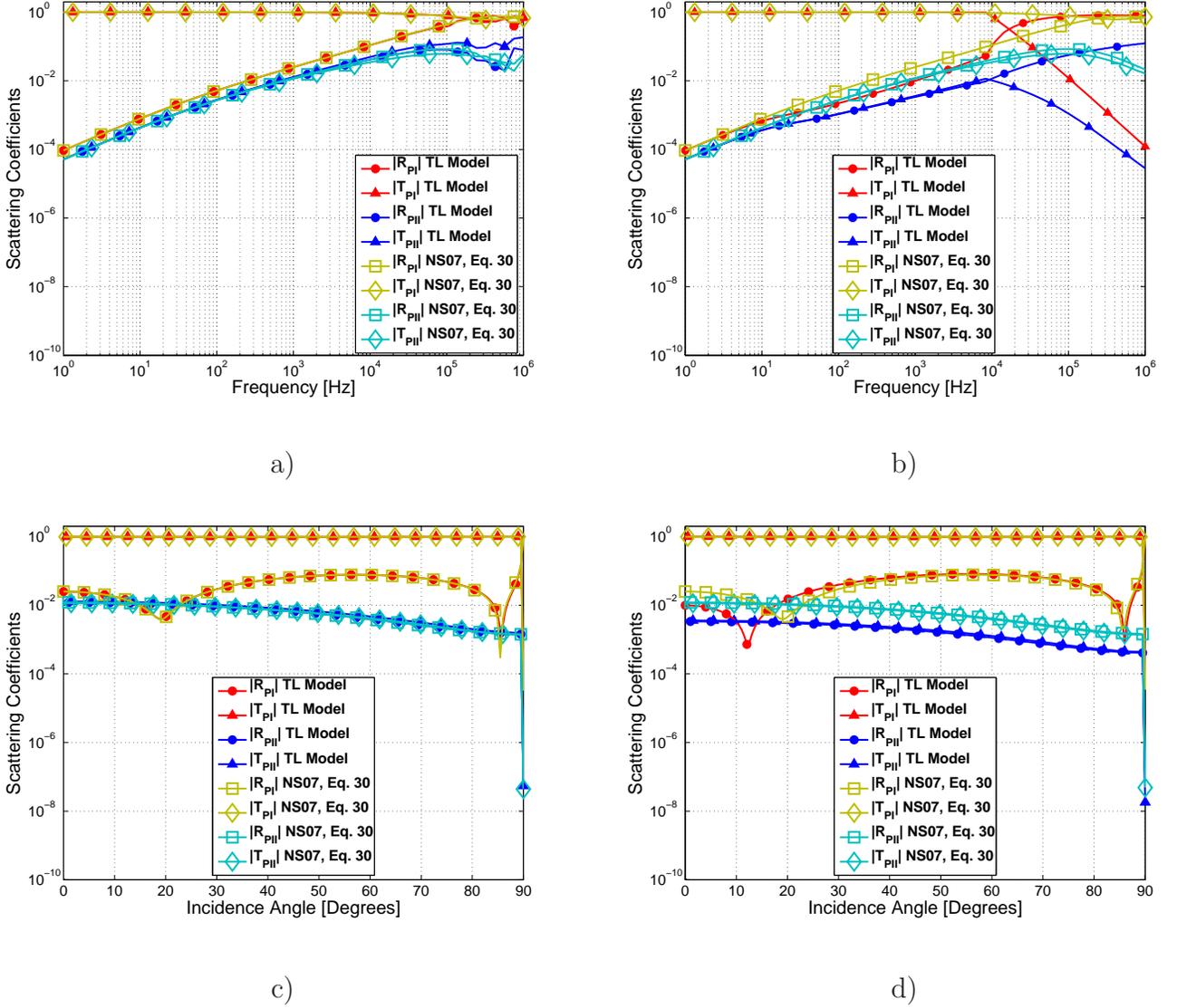


Figure 2: Absolute values of the reflection and transmission coefficients of type I (PI) and type II (PII) P-waves, for two values of permeability. The incident wave is a type I P-wave. The thickness of the fracture is $h = 0.001$ m. NS07 correspond to (31), and TL model to thin layer model. a) Normal incidence and permeability $\kappa_0 = 100$ D b) Normal incidence and permeability $\kappa_0 = 0.0001$ D. c) Frequency: 1000 Hz and permeability $\kappa_0 = 100$ D. d) Frequency: 1000 Hz and permeability $\kappa_0 = 0.0001$ D.

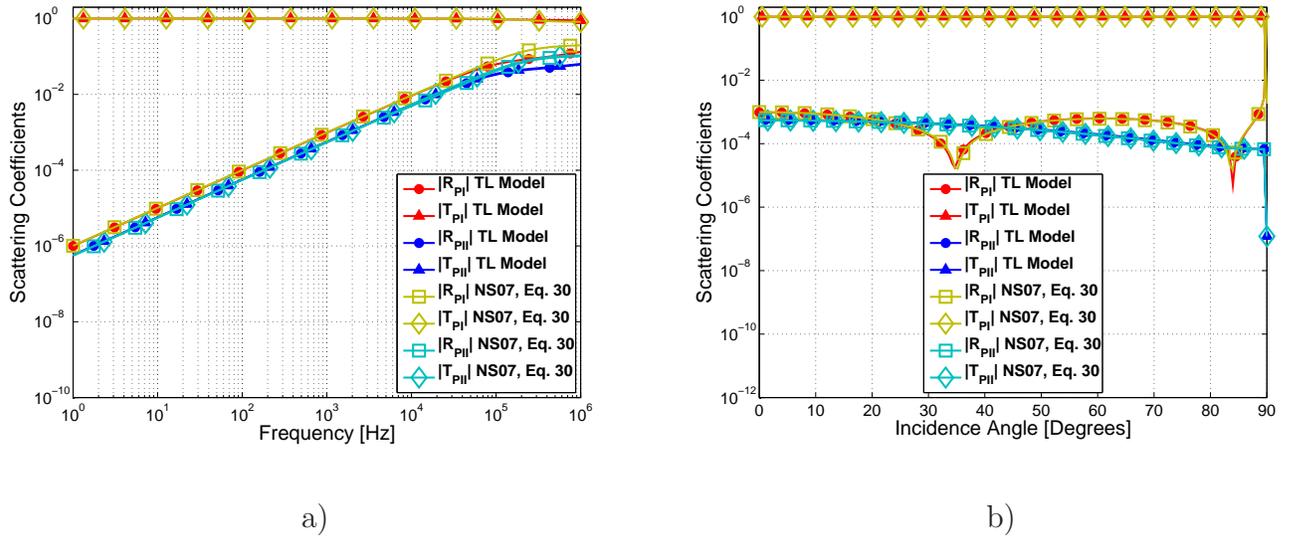


Figure 3: Absolute values of the reflection and transmission coefficients of type I (PI) and type (PII) P-waves. The incident wave is a type I P-wave. The thickness of fracture is $h = 0.00001$ m. NS07 correspond to (31), and TL model to thin layer model a) Normal incidence and permeability $\kappa_0 = 0.0001$ D. b) The frequency is 1000 Hz and the permeability is $\kappa_0 = 0.0001$ D.

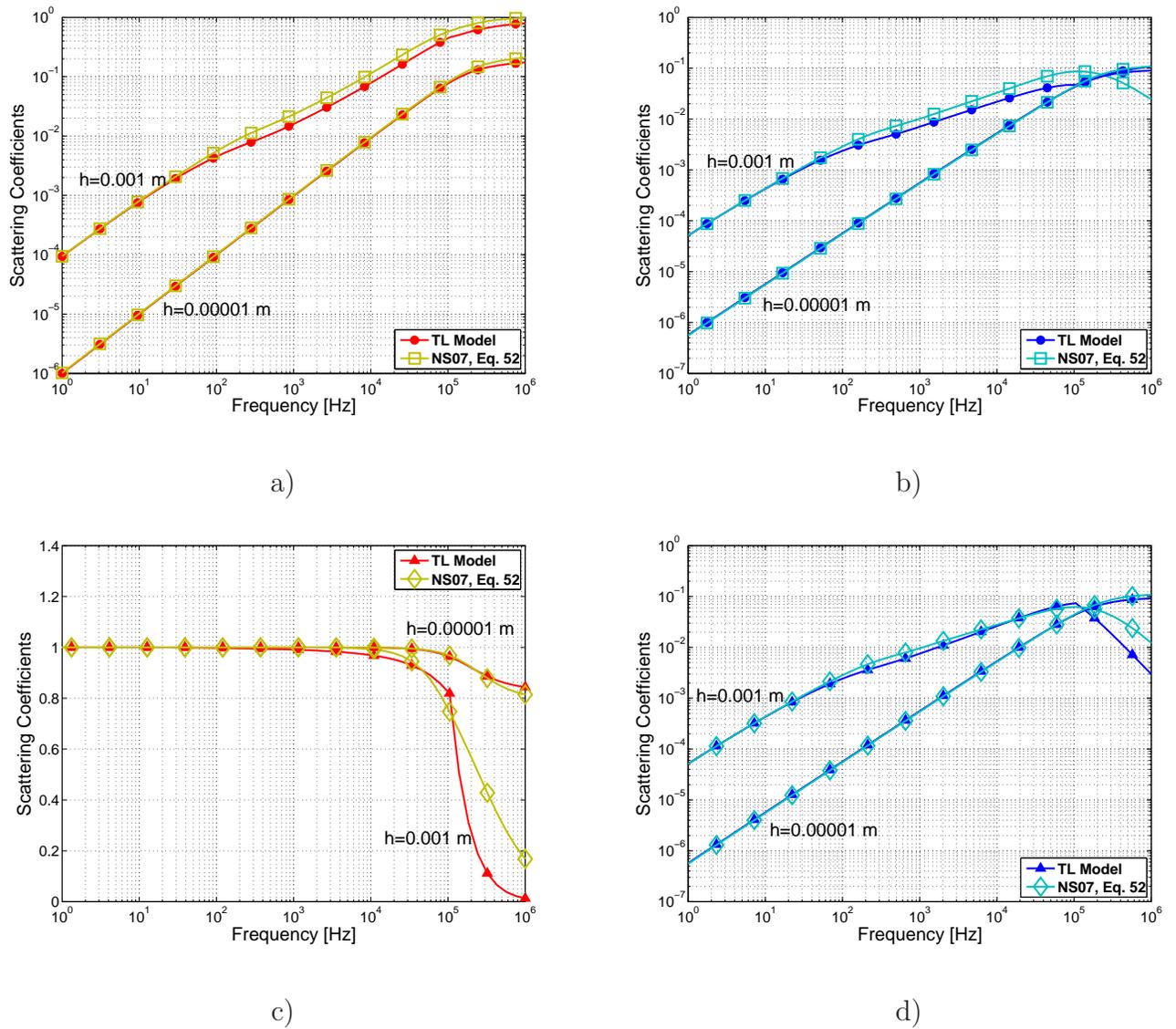
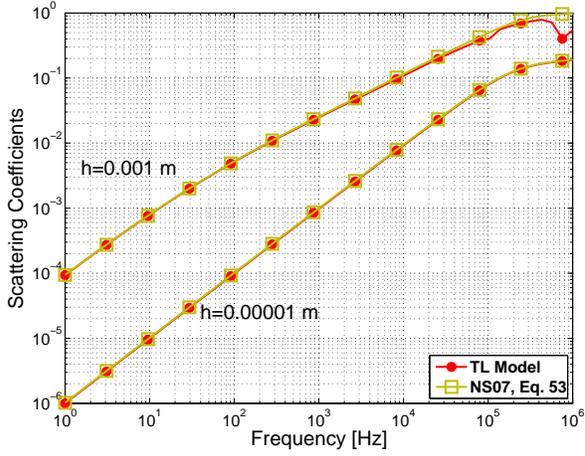
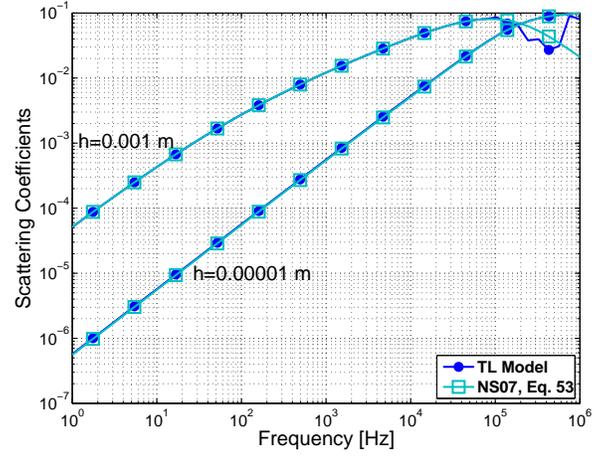


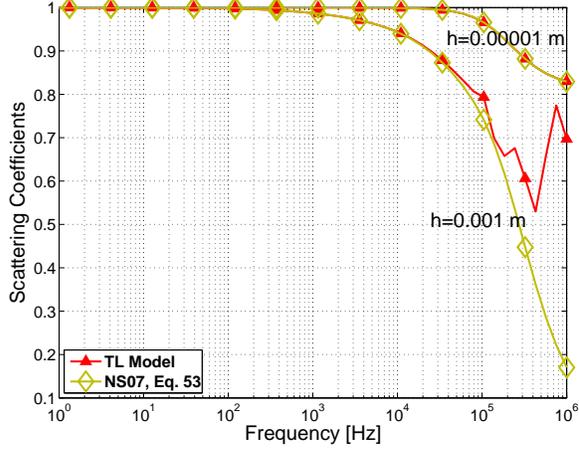
Figure 4: Reflection and transmission coefficients of type I and type II P-waves. The incident wave is a type I P-wave, permeability $\kappa_0 = 0.001$ D and h varies from 0.001 m to 0.00001 m. NS07 correspond to (31), and TL model to thin layer model. a) Absolute value of the reflection coefficient (type I wave); b) absolute value of the reflection coefficient (type II wave); c) absolute value of the transmission coefficient (type I wave) ; d) absolute value of the transmission coefficient (type I wave).



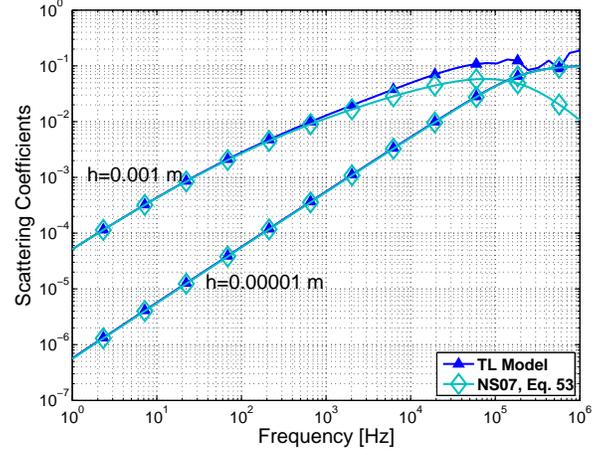
a)



b)

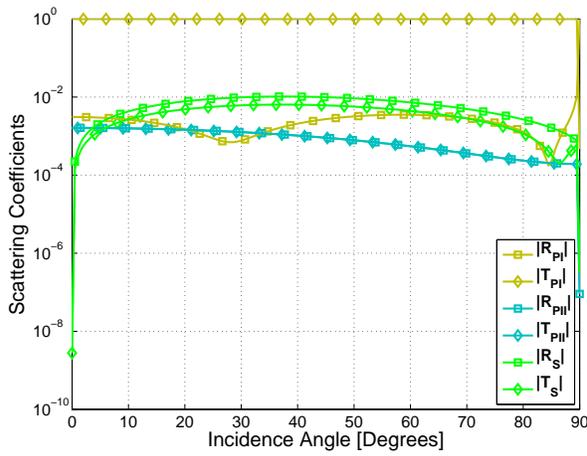


c)

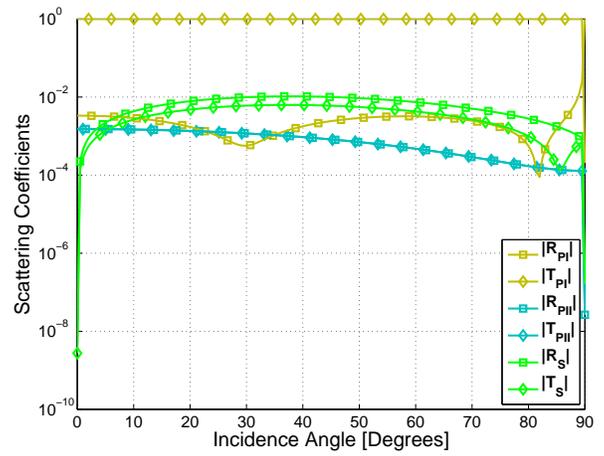


d)

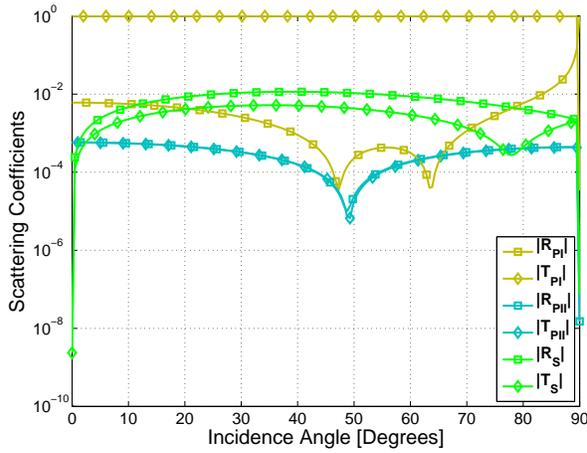
Figure 5: Reflection and transmission coefficients of type I (PI) and type II (PII) P-waves. The incident wave is a type I P-wave, permeability $\kappa_0 \rightarrow \infty$ and h varies from 0.001 m to 0.00001 m. NS07 correspond to (31), and TL model to thin layer model. a) Absolute value of the reflection coefficient (type I wave); b) absolute value of the reflection coefficient (type II wave); c) absolute value of the transmission coefficient (type I wave); e) absolute value of the transmission coefficient (type I wave).



a)



b)



c)

d)

Figure 6: Reflection and transmission coefficients of type I (PI), type II (PII) and shear waves as a function of the incidence angle. The coefficients were calculated using Equation 52 of (31). The incident wave is a type I P-wave of frequency 50 Hz, permeability $\kappa_0 = 100$ D and $h = 0.001$ m. a) Water in the fracture; b) oil in the fracture; c) gas in the fracture.